

Section A MASS WASTING

INTRODUCTION

This module summarizes the methods and results of a mass wasting assessment conducted on the Mendocino Redwood Company, LLC (MRC) ownership in the Howard Creek, Juan Creek, and Hardy Creek watershed, the Rockport Small Coastal Streams Watershed Analysis Unit (Rockport Coastal Streams WAU). California Planning Watersheds included in the Rockport Coastal Streams WAU include portions of the DeHaven Creek, Howard Creek, Juan Creek, and Hardy Creek Planning Watersheds. No landslides were detected on the relatively small area (~33 ac.) of MRC ownership within the DeHaven watershed during this assessment; this area was excluded from the analysis. This assessment was completed by Elias J. Steinbuck, PG 7538, and is part of a watershed analysis initiated by MRC that utilizes modified methodology adapted from procedures outlined in the Standard Methodology for Conducting Watershed Analysis (Version 4.0, Washington Forest Practices Board).

The principle objectives of this assessment are to:

- 1) Identify the types of mass wasting processes active in the basin.
- 2) Identify the link between mass wasting and forest management related activities.
- 3) Identify where the mass wasting processes are concentrated.
- 4) Partition the ownership into zones of relative mass wasting potential based on the likelihood of future mass wasting and sediment delivery to stream channels.

Additionally, the role of mass wasting sediment input to watercourses is examined. This information combined with the results of the Surface and Point Source Erosion module is used to construct a sediment input summary for the Rockport Coastal Streams WAU, contained in the Sediment Input Summary section of this watershed analysis.

The products of this report are: a landslide inventory map (Map A-1), a Terrain Stability Unit (TSU) map (Map A-2), and a mass wasting inventory database (Appendix A). The assembled information will enable forestland managers to make better forest management decisions to reduce management-induced risk of mass wasting. The mass wasting inventory will provide the information necessary to understand the spatial distribution, causal mechanisms, relative size, and timing of mass wasting processes active in the basin with reasonable confidence.

The Role of Mass Wasting in Watershed Dynamics

Mass wasting is a naturally occurring process, but can be accelerated by anthropogenic disturbances. Forest management practices can alter the natural frequency and magnitude of mass wasting events by changing the relative resisting and driving forces acting on a hillslope, altering soil and bedrock pore water pressures, and/or altering the effective cohesion of soil and bedrock (Sidle et.al., 1985). Increases in sediment yield due to mass wasting can disrupt the dynamic equilibrium of stream channels, resulting in a decline in the quality and quantity of amphibian and anadromous fish habitat, water quality, or stream ecology (Reeves et.al., 1995).

Mass wasting events are able to alter stream environments by increasing bed and suspended sediment loads, modifying the grain-size distribution of channel sediment, introducing woody debris, altering channel morphology by aggradation, damming and obstructing the channel, and in extreme cases scouring the channel to bedrock. Stream systems ultimately adjust

to major alterations downstream, as well as upstream of individual mass wasting events, however, the consequences may last for a long while.

In the Pacific Northwest where anadromous fish are present, mass wasting can have both beneficial and adverse effects on salmonid habitat. Beneficial effects include formation of new spawning, rearing, and over-wintering habitat due to addition of coarse gravels to the channel. The introduction of woody debris and boulders from landslides can increase cover and improve pool to riffle ratios. Adverse effects include filling of pools and scouring of riffles, blockage of fish access, disturbing side-channel rearing areas, and siltation of spawning gravels (Reeves et al., 1995). The magnitude of these effects are dependent on the frequency, location, and intensity of mass wasting events, as well as the sediment transporting capabilities of a particular stream. Beneficial and adverse effects typically occur simultaneously, and the relative relationship between the two will vary, even for individual events. Because of their greater stream powers, larger streams and rivers adjust to mass wasting perturbations faster than smaller streams.

BEDROCK STRUCTURE AND LITHOLOGY IN THE ROCKPORT COASTAL STREAMS WAU

The Rockport Coastal Streams WAU is underlain by bedrock of the Tertiary-Cretaceous Coastal Belt Franciscan, comprised predominately of interbedded sandstone and shale sequences with minor pebble conglomerate and greenstone (Kelly, 1983; Kelly, 1984). The Coastal Belt Franciscan is characterized by a relatively chaotic structure with shear zones, folds, and faults often juxtaposed with coherent sections of thin to massive sandstone and shale. Consistent with mass-flow type marine trench and trench-slope deposition, sedimentary structures are typically absent. Local alluvial deposits are present along the higher order channels within the Rockport Coastal Streams WAU.

The geomorphic expression of the Rockport Coastal Streams WAU is characterized by relatively short, steep basins that drain directly into the sea. No fault-rupture hazard zones, as depicted on the Alquist Priolo fault hazard maps, were identified in the region (DMG, 1997). Based on field reconnaissance, available geologic and hydrologic maps, and published literature, no regional indicators of adverse rock type, structure, or groundwater conditions were identified.

LANDSLIDE TYPES AND PROCESSES IN THE ROCKPORT COASTAL STREAMS WAU

Landslide features are widespread over the Rockport Coastal Streams WAU, owing to the relatively rapid down-cutting of the steep gradient creeks in response to global sea level fluctuations and regional uplift. The terminology used to describe landslides in this report closely follows the definitions of Cruden and Varnes (1996). This terminology is based on two nouns, the first describing the material that the landslide is composed of, and the second describing the type of movement. Landslides identified in the Rockport Coastal Streams WAU are discussed in detail below. For the purposes of this report, landslides are categorically separated into rapidly moving shallow-seated landslides and slowly moving deep-seated landslides, an essential distinction for forest management purposes.

Shallow-Seated Landslides

Debris slides, debris flows, and debris torrents are terms used throughout Mendocino Redwood Company's ownership to identify shallow-seated landslide processes. The material composition of debris slides, flows, or torrents is considered to be mainly colluvial soil with a significant proportion of coarse material; 20 to 80 percent of the particles are larger than 2 mm

(Cruden and Varnes, 1996). Shallow-seated slides generally move quickly downslope and commonly break apart during failure. Shallow-seated slides commonly occur along steep streamside slopes and in converging topography where colluvial materials accumulate and subsurface drainage concentrates. Susceptibility of a slope to fail by shallow-seated landslides is affected by slope steepness, saturation of soil, soil strength (friction angle and cohesion), and root strength (Selby, 1993). Due to the shallow depth and fact that debris slides, flows, or torrents involve the soil mantle, these are landslide types that can be significantly influenced by forest practices.

Debris slides are the most common landslide type observed in the WAU. The landslide mass typically fails along a surface of rupture or along relatively thin zones of intense shear strain located near the base of the soil profile. The landslide deposit commonly slides a distance beyond the toe of the surface of rupture and onto the ground surface below the failure; it generally does not slide more than the distance equal to the length of the failure scar. Landslides with deposits that traveled a longer distance below the failure scar would likely be defined as a debris flow or debris torrent. Debris slides commonly occur on steep planar slopes, convergent slopes, along forest roads and on steep slopes adjacent to watercourses. They usually fail by translational movement along an undulating or planar surface of failure. By definition debris slides do not continue downstream upon reaching a watercourse.

A debris flow is similar to a debris slide with the exception that the landslide mass continues to “flow” down the slope below the failure a considerable distance on top of the ground surface. A debris flow is characterized as a mobile, potentially very rapidly moving, slurry of soil, rock, vegetation, and water. High water content is needed for this process to occur. Debris flows generally occur on both steep, planar hillslopes and confined, convergent hillslopes. Often a failure will initiate as a debris slide, but will transform into a debris flow as it moves downslope.

Debris torrents have the greatest potential to destroy stream habitat and deliver large amounts of sediment. The main characteristic distinguishing a debris torrent is that the mass of failed soil and debris “torrents” downstream in a confined channel and erodes the bed and banks of the channel as it moves. As the debris torrent moves downslope and scours the channel, the liquefied landslide material generally increases in mass. Highly saturated soil or run-off in a channel is required for this process to occur. Debris torrents move rapidly and can potentially run down a channel for great distances. They typically initiate in headwall swales and torrent down intermittent watercourses. Often a failure will initiate as a debris slide, but will develop into a debris torrent upon reaching a channel. While actually a combination of two processes, these features were considered debris torrents.

Deep-Seated Landslides

Rockslides and earthflows are terms used throughout Mendocino Redwood Company’s ownership to identify deep-seated landslide processes. The failure dates of the deep-seated landslides could not be estimated with any confidence, they are likely to be of varying age with some potentially being over several thousand years old. Many of the deep-seated landslides are considered “dormant”, but the importance of identifying those lies in the fact that if reactivated, they have the potential to deliver large amounts of sediment and impair stream habitat. Accelerated or episodic movement is likely to have occurred over time in response to seismic shaking or high rainfall events.

Rockslides are deep-seated landslides with movement involving a relatively intact mass of rock and overlying earth materials. The failure plane is below the colluvial layer and involves the underlying bedrock. Mode of rock sliding generally is not strictly rotational or translational, but involves some component of each. Rotational slides typically fail along a concave surface, while translational slides typically fail on a planar or undulating surface of rupture. Rockslides

commonly create a flat, or back-tilted, bench below the crown of the scarp. A prominent bench is usually preserved over time and can be indicative of a rockslide. Rockslides fail in response to triggering mechanisms such as seismic shaking, adverse local structural geology, high rainfall, offloading or loading material on the slide, or channel incision (Wieczorek, 1996). The stream itself can be the cause of chronic movement, if it periodically undercuts the toe of a rockslide.

Earth flows are deep-seated landslides composed of fine-grained materials and soils derived from clay-bearing rocks. Earth flow materials typically consist of 80% or more of particles smaller than 2mm (Cruden and Varnes, 1996). Materials in an earth flow also commonly contain boulders, some very large, which move down slope in the clay matrix. Failure in earth flows is characterized by spatially differential rates of movement on discontinuous failure surfaces that are not preserved. The “flow” type of movement creates a landslide that can be very irregularly shaped. Some earth flow surfaces are dominantly grassland, while some are partially or completely forested. The areas of grassy vegetation are likely due to the inability of the unstable, clay-rich soils to support forest vegetation. The surface of an earth flow is characteristically hummocky with locally variable slope forms and relatively abundant gullies. The inherently weak materials within earth flows are not able to support steep slopes, therefore slope gradients are generally low to moderate. The rates of movement vary over time and can be accelerated by persistent high groundwater conditions. Timber harvesting can have the effect of increasing the amount of subsurface water, which can accelerate movement in an earth flow (Swanston et al, 1988).

Use of SHALSTAB by Mendocino Redwood Company for the Rockport Coastal Streams WAU

MRC uses SHALSTAB—a coupled steady state runoff infinite slope stability model—to assist with the mapping of the hazard potential of shallow-seated landslides (Dietrich and Montgomery, 1998). William Dietrich of the University of California (Berkeley) and David Montgomery of the University of Washington (Seattle) have published a validation study of the SHALSTAB model. Generally, they found that the SHALSTAB model correctly distinguishes areas more prone to shallow landslide instability. In mass wasting studies conducted in seven basins in northern California, they concluded that a log (q/T) threshold of less than -2.8 identifies the portion of the basin within which on average 57% of the shallow landslides mapped from aerial photographs are found. However, they also found that the performance of SHALSTAB depends strongly on the quality of the topographic data. The best readily available topographic data (10-m grid data from digitized USGS 7.5' quad maps) do not represent the fine scale topography that dictates the convergence of subsurface flow and the locations where shallow landslides are likely to occur. In our watershed analysis, we assess mass wasting hazards apart from SHALSTAB as well, using aerial photographs and field reconnaissance. However, we still use SHALSTAB output as one tool to assist with the interpretation of the landscape into terrain stability units.

METHODS

Landslide Inventory

The mass wasting assessment relies on an inventory of mass wasting features collected through the use of aerial photographs and field observations. MRC owned photographs from 2004 (color, 1:12,000), 2000 (color, 1:12,000), 1990 (color, 1:12,000), 1978 (color, 1:15,840), and 1972 (black and white, 1:15,840) were used, as were 1963 (black-and-white, 1:20,000), and 1952 (black-and-white, 1:20,000) photos on file at the Mendocino County Museum in Willits.

Data was collected regarding characteristics and measurements of the identified landslides. We acknowledge that some landslides may have been missed, particularly small ones that may be obscured by vegetation. A brief description of select parameters inventoried for each landslide observed in the field and during aerial photograph interpretation is presented in Figure A-1. A detailed discussion of these parameters follows.

Figure A-1. Description of Select Parameters used to describe Mass Wasting in the Mass Wasting Inventory.

- Slide Identification: Each landslide is assigned a unique identification number, a two letter code (see below) that denotes which planning watershed (PWS) the slide is located, and a number which indicates the USGS designated map section number the slide is mapped in.
 - Planning Watershed Codes:
 - RH – Hardy Creek
 - RJ – Juan Creek
 - RW – Howard Creek
- TSU # – Terrain Stability Unit in which landslide is located.
- Landslide Type:
 - DS – debris slide
 - DF – debris flow
 - DT – debris torrent
 - RS – rockslide
 - EF – earthflow
- Certainty: The certainty of identification is recorded.
 - D – Definite
 - P – Probable
 - Q – Questionable
- Physical Characteristics: Includes average length, width, depth, and volume of individual slides. Length of torrent, if present, is recorded as a comment.
- Sediment Routing: Denotes the type of stream the sediment was routed into.
 - P – Perennial
 - I – Intermittent or Ephemeral
 - N – No sediment delivered
- Sediment Delivery: Quantification of the relative percentage of the landslide that delivered to the stream.
- Slope: Percent slope angle is recorded for all shallow-seated landslides observed in the field.
- Age: Relative age of the observed slide is estimated.
 - N – new (<5 years old)
 - R – recent (5-10 years old)
 - O – old (>10 years old)
- Slope Form: Denotes morphology of the slope where the landslide originated
 - C – concave
 - D – divergent
 - P – planar
- Slide Location: Interpretation of the location where the landslide originated
 - H – Headwall Swale
 - S – Steep Streamside Slopes
 - I – Inner Gorge

N – Neither

- Road Association: Denotes the association of the landslide to land-use practices.
 - R – Road
 - S – Skid Trail
 - L – Landing
 - N – Neither
 - I – Indeterminate
- Contributing Area: Categorical description of the area interpreted to concentrate surface and/or subsurface flow to the point of failure for non-road related slide points.
 - S – Small, <0.5 acres
 - M – Medium, 0.5 – 3.0 acres
 - L – Large, >3.0 acres
- Aspect: Categorical description of the predominant cardinal direction the hillslope is facing for all slide points.
 - NE – Northeast, 0°-89°
 - NW – Northwest, 270°-359°
 - SE – Southeast aspect, 90°-179°
 - SW – Southwest aspect, 180°-269°
- Soil Type: County soil survey is used to attribute a soil type to each slide point. Soil types are grouped into similar grain size distributions based on the Unified Soil Classification System rating provided in the county survey.
 - C – Coarse, soils consisting of gravel-sand-silt mixture (GM-GC, USCS Class.)
 - F – Fine, soils consisting mainly of silt-clay (CL-ML, USCS Class.)
 - M – Mixed, soils with coarse and fine material (GC-CL)
- MRC Structure Class: 24 forest stand classes are used to describe the forest conditions across the MRC timberland. In this assessment this information is used to build a database of forest conditions upslope of recent (2001-2004 time period) non-road related failures. Structure classes are generated by classifying the following stand attributes:
 - Dominant Species
 - Dominant Diameter
 - Canopy Cover (%)
- Deep-seated landslides morphologic descriptions: toe, body, lateral scarps, and main scarp (see section below on Systematic Description of Deep-seated Landslide Features).

Landslides identified in the field and from aerial photograph observations are plotted on a landslide inventory map (Map A-1). All shallow-seated landslides are identified as a point plotted on the map at the interpreted head scarp of the failure. Deep-seated landslides are represented as a polygon representing the interpreted perimeter of the landslide body. Physical and geomorphic characteristics of all inventoried landslides are categorized in a database in Appendix A. Landslide dimensions and depths can be quite variable, therefore length, width, and depth values that are recorded are considered to be the average dimension of that feature. When converting landslide volumes to mass (tons), we assume a soil bulk density of 1.35 grams/cubic centimeter.

The certainty of landslide identification is assessed for each landslide. Three designations are used: definite, probable, and questionable. Definite means the landslide definitely exists. Probable means the landslide probably is there, but there is some doubt in the analyst's interpretation. Questionable means that the interpretation of the landslide identification may be inaccurate; the analyst has the least amount of confidence in the interpretation. Accuracy

in identifying landslides on aerial photographs is dependent on the size of the slide, scale of the photographs, thickness of canopy, and logging history. Landslides mapped in areas recently logged or through a thin canopy are identified with the highest level of confidence.

Characteristics of the particular aerial photographs used affects confidence in identifying landslides. For example, sun angle creates shadows which may obscure landslides, the print quality of some photo sets varies, and photographs taken at small scale makes identifying small landslides difficult. The landslide inventory results are considered a minimum estimate of sediment production. This is because landslides that were too small to identify on aerial photographs may have been missed, landslide surfaces could have reactivated in subsequent years and not been quantified, and secondary erosion by rills and gullies on slide surfaces is difficult to assess.

The technique employed to extrapolate a sediment volume delivery percentage to landslides not visited in the field relied on an average of those that were visited in the field. While this averaging technique is an oversimplification of actual on the ground sediment delivery measurements, it provides a means for estimating sediment delivery from the slides not visited in the field.

Landslides were classified based on the likelihood that a road associated land use practice was associated with the landslide. In this analysis, the effects of silvicultural techniques were not observed. The Rockport Coastal Streams WAU has been managed, recently and historically, for timber production. Therefore, it was determined that the effect of silvicultural practices was too difficult to confidently assign to landslides. There have been too many different silvicultural activities over time for reasonable confidence in a landslide evaluation based on silviculture. The land use practices that were assigned to landslides were associations with roads, skid trails, or landings. It was assumed that a landslide adjacent to a road, skid trail, or landing was triggered either directly or indirectly by that land use practice. If a landslide appeared to be influenced by more than one land use practice, the more causative one was noted. If a cutslope failure did not cross the road prism, it was assumed that the failure would remain perched on the road, landing, or skid trail and would not deliver to a watercourse. Some surface erosion could result from a cutslope failure and is assumed to be addressed in the road surface erosion estimates (Surface and Fluvial Erosion Module).

Sediment Input from Shallow-Seated Landslides

The overall time period used for mass wasting interpretation and sediment budget analysis is sixty-seven years. Sediment input to stream channels by mass wasting is quantified for seven time periods (1942-1952, 1953-1963, 1964-1972, 1973-1978, 1979-1990, 1991-2000, 2001-2004). The evaluation assumes that approximately the last 10 years of mass wasting can be observed in the aerial photograph. This is due to landslide surfaces revegetating quickly, making small mass wasting features older than about 10 years difficult to see. We acknowledge that we have likely missed an unknown quantity of small mass wasting events during the aerial photograph interpretation. However, we assume we have captured the majority of the larger mass wasting events in this analysis.

Sediment delivery estimates from mapped shallow-seated landslides were used to produce the total mass wasting sediment input. In order to extrapolate depth to the shallow-seated landslides not visited in the field, an average was taken from the measured depths of landslides visited in the field. Field measurements revealed a similar distribution of depths for management associated (which includes roads, skid trails, and landings), and non-management associated shallow-seated landslides. Therefore, the shallow-seated landslides not verified in the field were assigned the average depth from field verified landslides. In order to extrapolate sediment delivery percentage to landslides not verified in the field, an average was taken from the estimated delivery percentage of field verified landslides.

Delivery statistics were not calculated for deep-seated landslides; however, some of the sediment delivery from shallow-seated landslides is the result of conditions created by deep-seated landslides. For example, a deep-seated failure could result in a debris slide or torrent, which could deliver sediment. Furthermore, over-steepened scarps or toes of deep-seated landslides may have shallow failures associated with them. These types of sediment delivery from shallow-seated landslides associated with deep-seated landslides are accounted for in the delivery estimates.

Sediment Input from Deep-Seated Landslides

Large, active, deep-seated landslides can potentially deliver large volumes of sediment. Delivery generally occurs over long time periods compared to shallow-seated landslides, with movement delivering earth materials into the channel, resulting in an increased sediment load downstream of the failure. Actual delivery can occur by over-steepening of the toe of the slide and subsequent failure into the creek, or by the slide pushing out into the creek. It is very important not to confuse normal stream bank erosion at the toe of a slide as an indicator of movement of that slide. Before making such a connection, the slide surface should be carefully explored for evidence of significant movement, such as wide ground cracks. Sediment delivery could also occur in a catastrophic manner. In such a situation, large portions of the landslide essentially fail and move into the watercourse “instantaneously”. These types of deep-seated failures are relatively rare on MRC property and usually occur in response to unusual storm events or seismic ground shaking.

Movement of deep-seated landslides has definitely resulted in some sediment delivery in the Rockport Coastal Streams WAU. Quantification of the sediment delivery from deep-seated landslides was not determined in this watershed analysis. Factors such as rate of movement, or depth to the slide plane, are difficult to determine without subsurface geotechnical investigations that were not conducted in this analysis. Sediment delivery to watercourses from deep-seated landslides (landslides typically ≥ 10 feet thick) can occur by several processes. Such processes can include surface erosion and shallow-or deep-seated movement of a portion or all of the deep-seated landslide deposit.

The ground surface of a deep-seated landslide, like any other hillside surface, is subject to surface erosion processes such as rain drop impact, sheet wash (overland flow), and gully/rill erosion. Under these conditions the sediment delivery from surficial processes is assumed the same as adjacent hillside slopes not underlain by landslide deposits. The materials within the landslide are disturbed and can be arguably somewhat weaker. However, once a soil has developed, the fact that the slope is underlain by a deep-seated landslide should make little difference regarding sediment delivery generated by erosional processes that act at the ground surface. Although fresh, unprotected surfaces that develop in response to recent or active movement could become a source of sediment until the bare surface becomes covered with leaf litter, re-vegetated, or soils developed.

Clearly, movement of a portion or all of a deep-seated landslide can result in delivery of sediment to a watercourse. This determination is made by exploring for any evidence of movement. However, movement would need to be on slopes immediately adjacent to or in close proximity to a watercourse and of sufficient magnitude to push the toe of the slide into the watercourse. A deep-seated slide that toes out on a slope far from a creek or moves only a short distance downslope will generally deliver little to a watercourse. It is also important to realize that often only a portion of a deep-seated slide may become active, though the portion could be quite variable in size. Ground cracking at the head of a large, deep-seated landslide does not necessarily equate to immediate sediment delivery at the toe of the landslide. Movement of large deep-seated landslides can create void spaces within the slide mass. Though movement can be clearly indicated by the ground cracks, many times the toe may not respond or show indications

of movement until some of the void space is “closed up”. This would be particularly true in the case of very large deep-seated landslides that exhibit ground cracks that are only a few inches to a couple of feet wide. Compared to the entire length of the slide, the amount of movement implied by the ground crack could be very small. This combined with the closing up or “bulking up” of the slide, would not generate much movement, if any, at the toe of the slide. Significant movement, represented by large wide ground cracks, would need to occur to result in significant movement and sediment delivery at the toe of the slide.

Systematic Description of Deep-seated Landslide Features

The characteristics of deep-seated landslides received less attention in the landslide inventory than shallow-seated landslides mainly due to the fact that subsurface analyses would have to be conducted to estimate attributes such as depth, volume, failure date, current activity, and sediment delivery. Subsurface investigation was beyond the scope of this report. Few of the mapped deep-seated landslides were observed to have recent movement associated with them, mainly due to oversteepening of the slope at the toe or scarp. Further assessment of deep-seated landslides will occur on a site-by-site basis in the Rockport Coastal Streams WAU, likely during timber harvest plan preparation and review.

Deep-seated landslides were mainly interpreted by reconnaissance techniques (aerial photograph interpretation complemented by limited field observations). Reconnaissance mapping criteria consist of observations of four morphologic features of deep seated landslides – toe, internal morphology, lateral flanks, main scarp, and vegetation (after McCalpin 1984 as presented by Keaton and DeGraff, 1996, p. 186, Table 9-1). The mapping and classification criteria for each feature are presented in detail below.

Aerial photo interpretation of deep seated landslide features in the Rockport Coastal Streams WAU suggests that the first three morphologic features above are the most useful for inferring the presence of deep-seated landslides. The presence of tension cracks and/or sharply defined and topographically offset scarps are probably a more accurate indicator of recent or active landslide movement. These features, however, are rarely visible on aerial photos.

Sets of five descriptions have been developed to classify each deep-seated landslide morphologic feature or vegetation influence. The five descriptions are ranked in descending order from characteristics more typical of active landslides to dormant to relict landslides. One description should characterize the feature most accurately. Nevertheless, some overlap between classifications is neither unusual nor unexpected. We recognize that some deep-seated landslides may lack evidence with respect to one or more of the observable features, but show strong evidence of another feature. If there is no expression of a particular geomorphic feature (e.g. lateral flanks), the classification of that feature is considered “undetermined”. If a deep-seated landslide is associated with other deep-seated landslides, it may also be classified as a landslide complex.

In addition to the classification criteria specific to the deep-seated landslide features, more general classification of the strength of the interpretation of the deep-seated landslide is conducted. Some landslides are obscured by vegetation to varying degrees, with areas that are clearly visible and areas that are poorly visible. In addition, weathering and erosion processes may also obscure geomorphic features over time. The quality of different aerial photograph sets varies and can sometimes make interpretations difficult. Owing to these circumstances, each inferred deep-seated landslide feature is classified according to the strength of the evidence as definite, probable or questionable as defined with respect to interpretation of shallow landslides.

At the project scale (THP development and planning), field observations of deep-seated landslide morphology and other indicators by qualified professionals are expected to be used to reduce uncertainty of interpretation inherent in reconnaissance mapping. Field criteria for mapping deep-seated landslides and assessment of activity are presented elsewhere.

Deep Seated Landslide Morphologic Classification Criteria:**I. Toe Activity**

1. Steep streamside slopes with extensive unvegetated to sparsely vegetated debris slide scars. Debris slides occur on both sides of stream channel, but more prominently on side containing the deep-seated landslide. Stream channel in toe region may contain coarser sediment than adjacent channel. Stream channel may be pushed out by toe. Toe may be eroding, sharp topography/geomorphology.
2. Steep streamside slopes with few unvegetated to sparsely vegetated debris slide scars. Debris slides generally are distinguishable only on streamside slope containing the deep-seated landslide. Stream channel may be pushed out by toe. Sharp edges becoming subdued.
3. Steep streamside slopes that are predominantly vegetated with little to no debris slide activity. Topography/geomorphology subdued.
4. Gently sloping stream banks that are vegetated and lack debris slide activity. Topography/geomorphology very subdued.
5. Undetermined

II. Internal Morphology

1. Multiple, well defined scarps and associated angular benches. Some benches may be rotated against scarps so that their surfaces slope back into the hill causing ponded water, which can be identified by different vegetation than adjacent areas. Hummocky topography with ground cracks. Jack-strawed trees may be present. No drainage to chaotic drainage/disrupted drainage.
2. Hummocky topography with identifiable scarps and benches, but those features have been smoothed. Undrained to drained but somewhat subdued depressions may exist. Poorly established drainage.
3. Slight benches can be identified, but are subtle and not prominent. Undrained depressions have since been drained. Moderately developed drainage to established drainage but not strongly incised. Subdued depressions but are being filled.
4. Smooth topography. Body of slide typically appears to have failed as one large coherent mass, rather than broken and fragmented. Developed drainage well established, incised. Essentially only large undrained depressions preserved and would be very subdued. Could have standing water. May appear as amphitheater slope where slide deposit is mostly or all removed.
5. Undetermined

III. Lateral Flanks

1. Sharp, well defined. Debris slides on lateral scarps fail onto body of slide. Gullies/drainage may begin to form at boundary between lateral scarps and sides of slide deposit. Bare spots are common or partially unvegetated.
2. Sharp to somewhat subdued, rounded, and essentially continuous, might have small breaks; gullies/drainage may be developing down lateral edges of slide body. May have debris slide activity, but less prominent. Few bare spots.

3. Smooth, subdued, but can be discontinuous and vegetated. Drainage may begin to develop along boundary between lateral scarp and slide body. Tributaries to drainage extend onto body of slide.
4. Subtle, well subdued to indistinguishable, discontinuous. Vegetation is identical to adjacent areas. Watercourses could be well incised, may have developed along boundary between lateral scarp and slide body. Tributaries to drainage developed on slide body.
5. Undetermined

IV. Main Scarp

1. Sharp, continuous geomorphic expression, usually arcuate break in slope with bare spots to unvegetated; often has debris slide activity.
2. Distinct, essentially continuous break in slope that may be smooth to slightly subdued in parts and sharp in others, apparent lack of debris slide activity. Bare spots may exist, but are few.
3. Smooth, subdued, less distinct break in slope with generally similar vegetation relative to adjacent areas. Bare spots are essentially non-existent.
4. Very subtle to subdued, well vegetated, can be discontinuous and deeply incised, dissected; feature may be indistinct.
5. Undetermined

V. Vegetation

1. Less dense vegetation than adjacent areas. Recent slide scarps and deposits leave many bare areas. Bare areas also due to lack of vegetative ability to root in unstable soils. Open canopy, may have jack-strawed trees; can have large openings.
2. Bare areas exist with some regrowth. Regrowth or successional patterns related to scarps and deposits. May have some openings in canopy or young broad-leaf vegetation with similar age.
3. Subtle differences from surrounding areas. Slightly less dense and different type vegetation. Essentially closed canopy; may have moderately aged to old trees.
4. Same size, type, and density as surrounding areas.
5. Undetermined

Terrain Stability Units

Terrain Stability Units (TSUs) are delineated by partitioning the landscape into zones characterized by similar geomorphic attributes, shallow-seated landslide potential, and sediment delivery to stream channels. A combination of aerial photograph interpretation, field investigation, and SHALSTAB output were utilized to delineate TSUs. The TSU designations for the Rockport Coastal Streams WAU are only meant to be general characterizations of similar geomorphic and terrain characteristics related to shallow seated landslides. Deep-seated landslides are also shown on the TSU map (Map A-2). The deep-seated landslides have been included to provide land managers with supplemental information to guide evaluation of harvest planning and subsequent needs for geologic review. The landscape and geomorphic setting in the Rockport Coastal Streams WAU is certainly more complex than generalized TSUs delineated for this evaluation. The TSUs are only meant to be a starting point for gauging the need for site-specific field assessments.

The delineation of each TSU described is based on landforms present, the mass wasting processes, sensitivity to forest practices, mass wasting hazard, delivery potential, and forest

management related trigger mechanisms for shallow seated landslides. The landform section of the TSU description defines the terrain found within the TSU. The mass wasting process section is a summary of landslide types found in the TSU. Sensitivity to forest practice and mass wasting hazard is, in part, a subjective call by the analyst based on the relative landslide hazard and influence of forest practices. Delivery potential is based on proximity of TSU to watercourses and the likelihood of mass wasting in the unit to reach a watercourse. The hazard potential is based on a combination of the mass wasting hazard and delivery potential (Table A-1). The trigger mechanisms are a list of forest management practices that may have the potential to create mass wasting in the TSU.

Table A-1. Ratings for Potential Hazard of Delivery of Debris and Sediment to Streams by Mass Wasting (L= low hazard, M= moderate hazard, H = high hazard) (from Version 4.0, Washington Forest Practices Board, 1995).

		Mass Wasting Potential		
		Low	Moderate	High
Delivery Potential	Low	L	L	M
	Moderate	L	M	H
	High	L	M	H

RESULTS

Mass Wasting Inventory

A Landslide Inventory Data Sheet (Appendix A) was used to record attributes associated with each landslide. The spatial distribution and location of landslides is shown on Map A-1.

A total of 412 shallow-seated landslides (debris slides, torrents, or flows) were identified and characterized in the Rockport Coastal Streams WAU. A total of 41 deep-seated landslides (rockslides and earthflows) were mapped in the Rockport Coastal Streams WAU. A considerable effort was made to field verify as many landslides as possible to insure greater confidence in the results. Approximately 21% (86/412) of the identified shallow-seated landslides were field verified. From this level of field observations, extrapolation of landslide depth and sediment delivery is assumed to be performed with a reasonable level of confidence.

The temporal distribution of the 412 shallow-seated landslides observed in the Rockport Coastal Streams WAU is listed in Table A-2. The distribution by landslide type is shown in Table A-3.

Table A-2. Shallow-Seated Landslide Summary for Rockport Coastal Streams WAU by Time Periods.

Planning Watershed	1943 - 1952	1953 - 1963	1964 - 1972	1973 - 1978	1979 - 1990	1991 - 2000	2001 - 2004
Hardy Creek	28	12	12	16	18	10	3
Juan Creek	20	7	42	111	33	19	7
Howard Creek	7	9	9	16	16	17	0
RP Coastal WAU	55	28	63	143	67	46	10

Table A-3. Landslide Summary by Type and Planning Watershed for Rockport Coastal Streams WAU.

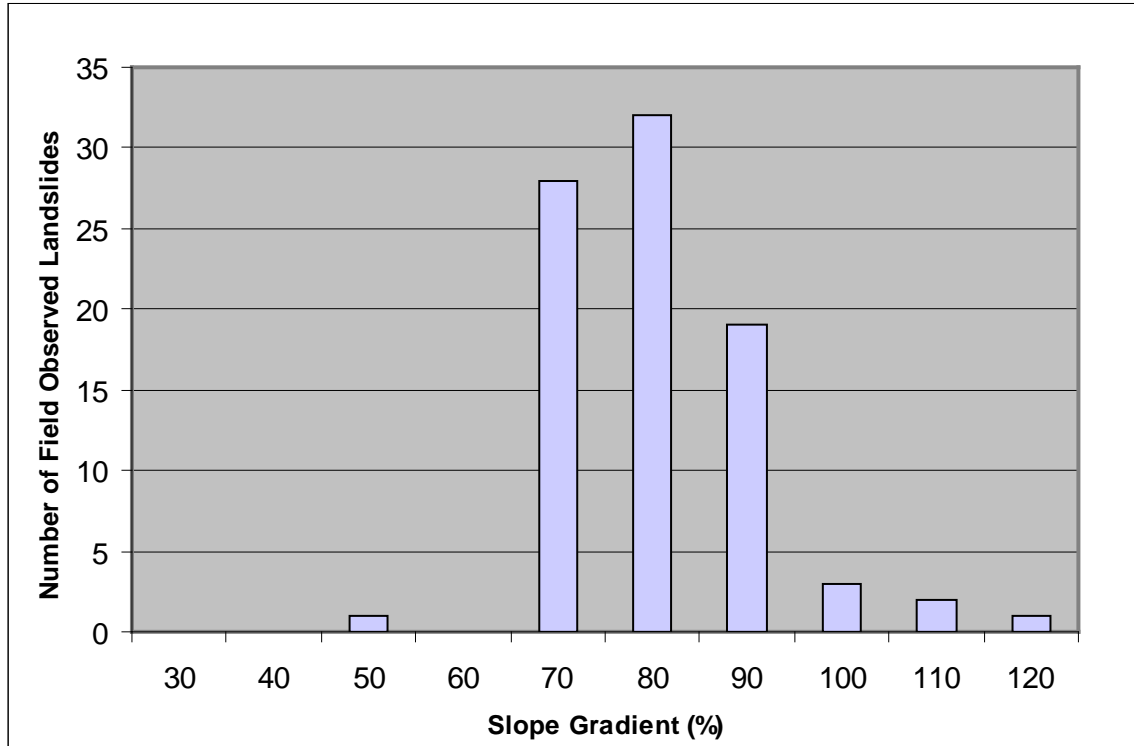
Planning Watershed	Debris Slides	Debris Flows	Debris Torrents	Rock-slides	Earth-flows	Total	Road ^a Assoc.
Hardy Creek	91	5	3	5	0	0	43
Juan Creek	231	3	5	20	0	0	208
Howard Creek	68	2	4	16	0	0	50
Rockport Coastal Streams WAU	390	10	12	41	0	453	301

a – Includes roads, skid trails, and landings

The majority of the landslides observed in the Rockport Coastal Streams WAU are debris slides. Of the 412 shallow-seated landslides in the Rockport Coastal Streams WAU, 301 are determined to be road associated (includes roads, skid trails, or landings). This is approximately 73% of the total number of shallow-seated landslides. There were 22 debris torrents and flows observed in the Rockport Coastal Streams WAU. This is approximately 5% of the total shallow-seated landslides observed in the Rockport Coastal Streams WAU.

Of the 86 field observed shallow-seated landslides across the Rockport Coastal Streams WAU, 99% (85/86) were initiated on slopes of 70% gradient or higher (Chart A-1).

Chart A-1. Slope Gradient Histogram for Shallow-Seated Landslides Occurring on MRC Ownership in the Rockport Coastal Streams WAU.



A majority of inventoried landslides originated in convergent topography (231/412, or 56%) where subsurface water tends to concentrate, or on steep, planar topography (179/412, or 43%), where sub-surface water can be concentrated at the base of slopes, in localized topographic depressions, or by local geologic structure. Few landslides originated in divergent topography (2/412, or <1%), where subsurface water is typically routed to the sides of ridges (Chart A-2).

A majority of the inventoried landslides were discovered along steep streamside slopes (266/412, or 65%), with fewer found in headwall swales (58/412, or 14%) and inner gorge slopes (11/412, or 3%) observed along the outside edge of meander bends. A significant portion (77/412, or 19%) of the inventoried landslides were observed on open slopes away from any inner gorge, steep streamside slopes, or headwall swales, however, a majority of these slides originated in fill material along the outside edge of roads and skid trails (Chart A-3). Such observations were, in part, the basis for the delineation of the WAU into Terrain Stability Units.

Chart A-2. Slope Morphology Summary for Shallow-Seated Landslides Occurring on MRC Ownership in the Rockport Coastal Streams WAU.

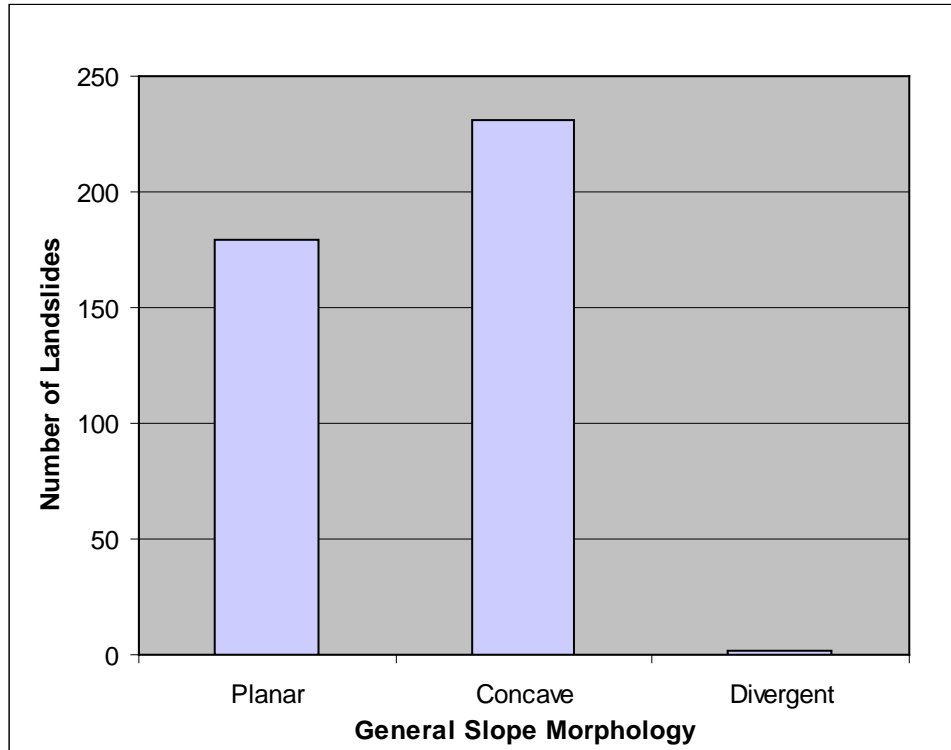
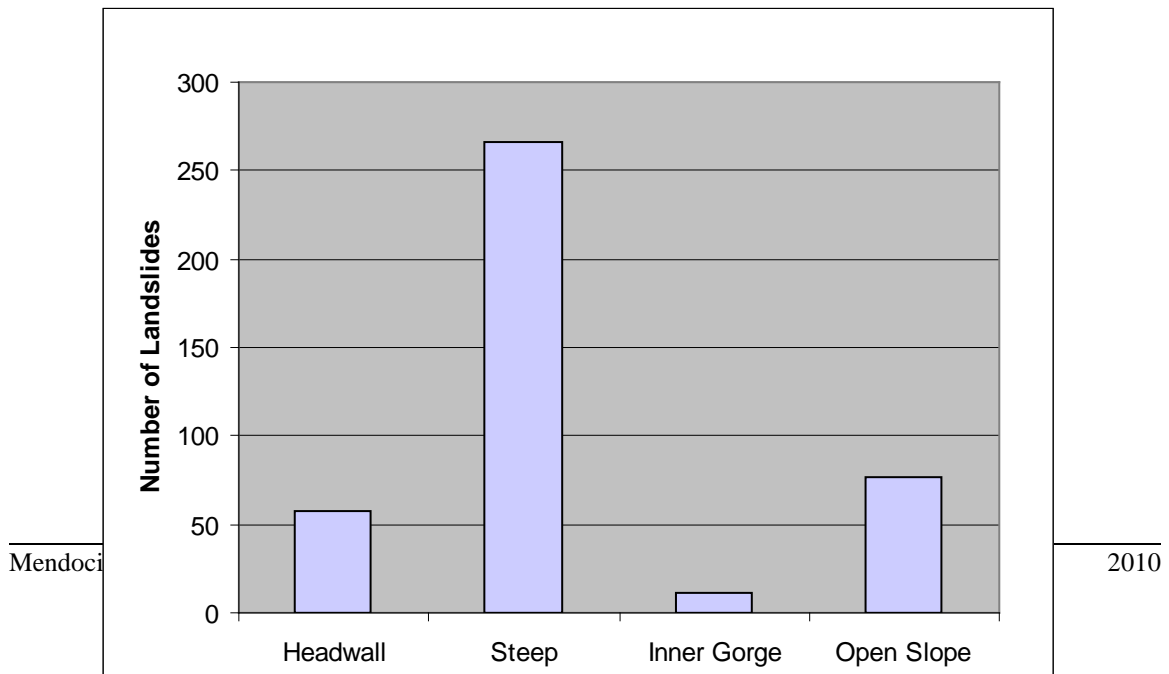


Chart A-3. Slide Location Summary for Shallow-Seated Landslides Occurring on MRC Ownership in the Rockport Coastal Streams WAU.



Terrain Stability Units

The landscape was partitioned into seven Terrain Stability Units representing general areas of similar geomorphology, landslide processes, and sediment delivery potential for shallow-seated landslides (Map A-2). The units are to be used by forest managers to assist in making decisions that will minimize future mass wasting sediment input to watercourses. The delineation for the TSUs was based on qualitative observations and interpretations from aerial photographs, field evaluation, and SHALSTAB output. Deep-seated landslides are also shown on the TSU map (Map A-2). The deep-seated landslides have been included to provide land managers with supplemental information to guide evaluation of harvest planning and subsequent needs for geologic review.

Shallow-seated landslide characteristics considered in determination of map units are size, frequency, delivery to watercourses, and spatial distribution. Hillslope characteristics considered are slope form (convergence, divergence, planar), slope gradient, relative magnitude of stream incision, and overall geomorphology. The range of slope gradients was determined from USGS 1:24,000 topographic maps and field observations. Hillslope and landslide morphology vary within each individual TSU and the boundaries are not exact. This evaluation is not intended to be a substitute for site-specific field assessments. Site-specific field assessments will still be required in TSUs and at deep-seated landslides or specific areas of some TSUs to assess the risk and likelihood of mass wasting impacts from a proposed management action. The TSUs are compiled on the entitled Terrain Stability Unit Map (Map A-2).

TSU Number:	1
Description:	Inner Gorge or Steep Streamside Slopes adjacent to Low Gradient Watercourses
Materials:	Shallow soils formed on weathered marine sedimentary rocks. Maybe composed of toe sediment of deep-seated landslide deposit.
Landform:	Characterized by steep streamside slopes or inner gorge topography along low gradient watercourses (typically less than 6-7%). An inner gorge is a geomorphic feature created from down cutting of the stream, generally in response to tectonic uplift. Inner gorge slopes extend from either one or both sides of the stream channel to the first break in slope. Inner gorge slope gradients typically exceed 70%, although slopes with lower inclination are locally present. Inner gorge slopes commonly contain areas of multiple, coalescing shallow seated landslide scars of varying age. Steep streamside slopes are characterized by their lack of a prominent break in slope. Slopes are generally planar in form with slope gradients typically exceeding 70%. The upper extent of TSU 1 is variable. Where there is not a break in slope, the unit may extend 300 feet upslope (based on the range of lengths of landslides observed, 20-300 feet). Landslides in this unit generally deposit sediment directly into Class I and II streams. Small areas of incised terraces may be locally present.
Slope:	Typically >70 %, (mean slope of observed mass wasting events is 83%, range is 70%-110%)
Total Area:	1,037 acres; 10% of the total WAU area.
MW Processes:	<p>83 <i>road-associated landslides</i></p> <ul style="list-style-type: none"> • 80 Debris slides • 1 Debris flows • 2 Debris torrents <p>69 <i>non-road associated landslides</i></p> <ul style="list-style-type: none"> • 66 Debris slides • 2 Debris flows • 1 Debris torrents
Non Road-related Landslide Density:	0.07 landslides per acre for the past 62 years.
Forest Practices Sensitivity:	High sensitivity to road construction due to proximity to watercourses, high sensitivity to harvesting and forest management practices due to steep slopes with localized colluvial or alluvial soil deposits adjacent to watercourses.
Mass Wasting Potential:	High localized potential for landslides in both unmanaged and managed conditions.

Delivery Potential: High

Delivery Criteria

Used: Steep slopes adjacent to stream channels, a majority of the observed landslides delivered sediment into streams.

Hazard-Potential

Rating: **High**

Forest Management

Related Trigger

Mechanisms:

- Sidecast fill material placed on steep slopes can initiate debris slides or flows in this unit.
- Concentrated drainage from roads onto unstable areas can initiate debris slides or flows in this unit.
- Poorly sized culvert or excessive debris at watercourse crossings can initiate failure of the fill material creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Sidecast fill material created from skid trail construction placed on steep slopes can initiate debris slides or flows in this unit.
- Concentrated drainage from skid trails onto unstable areas can initiate debris slides or flows in this unit.
- Cut-slope of skid trails can remove support of slope creating debris slides, torrents or flows in this unit.
- Cut-slope of skid trails can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Concentrated drainage from roads can increase groundwater, accelerating movement of rockslides or earth flows and over-steepening TSU 1 slopes.
- Removal of vegetation from these slopes can result in loss of evapotranspiration and thus increase pore water pressures that could initiate slope failure in this unit.
- Post timber harvest root decay of hardwood or non-redwood conifer species can be a contributing factor in the initiation of debris slides, torrents or flows in this unit.

Confidence:

High confidence for susceptibility of landslides and sediment delivery in this unit. Moderate confidence in placement of the unit boundary. This unit is locally variable and exact boundaries are best determined during field observations. Within this unit there are likely areas of low gradient slopes that are less susceptible to mass wasting.

TSU Number:	2
Description:	Inner gorge or Steep Streamside Slopes adjacent to high gradient intermittent or ephemeral watercourses.
Materials:	Shallow soils formed from weathered marine sedimentary rocks with localized areas of thin to thick colluvial deposits.
Landforms:	Characterized by steep streamside slopes or inner gorge topography along low gradient watercourses (typically greater than 6-7%). An inner gorge is a geomorphic feature created from down cutting of the stream, generally in response to tectonic uplift. Inner gorge slopes extend from either one or both sides of the stream channel to the first break in slope. Inner gorge slope gradients typically exceed 70%, although slopes with lower inclination are locally present. Inner gorge slopes commonly contain areas of multiple, coalescing shallow seated landslide scars of varying age. Steep streamside slopes are characterized by their lack of a prominent break in slope. Slopes are generally planar in form with slope gradients typically exceeding 70%. The upper extent of TSU 2 is variable. Where there is not a break in slope, the unit may extend 300 feet upslope (based on the range of lengths of landslides observed, 25-300 feet). Landslides in this unit generally deposit sediment directly into Class I and II streams.
Slope:	Typically >70% (mean slope of observed mass wasting events is 82%, range is 55%-120%).
Total Area:	1,059 acres; 10% of total WAU area
MW Processes:	<p>104 <i>road-associated landslides</i></p> <ul style="list-style-type: none"> • 99 Debris slides • 1 Debris flows • 4 Debris torrents <p>22 <i>non-road associated landslides</i></p> <ul style="list-style-type: none"> • 21 Debris slides • 0 Debris flows • 1 Debris torrent
Non Road-related Landslide Density:	0.02 landslides per acre for the past 62 years.
Forest Practices Sensitivity:	High sensitivity to roads due to steep slopes adjacent to watercourses, high to moderate sensitivity to harvesting and forest management due to steep slopes next to watercourses. Localized areas of steeper and/or convergent slopes may have an even higher sensitivity to forest practices.

Mass Wasting

Potential: High in both unmanaged and managed conditions due to the steep morphology of the slope.

Delivery Potential: High

Delivery Criteria

Used: Steep slopes adjacent to stream channels, a majority of the observed landslides delivered sediment into streams.

Hazard-Potential

Rating: **High**

Forest Management

Related Trigger

Mechanisms:

- Sidecast fill material placed on steep slopes can initiate debris slides, torrents or flows in this unit.
- Concentrated drainage from roads onto unstable areas can initiate debris slides, torrents or flows in this unit.
- Poorly sized culvert or excessive debris at watercourse crossings can initiate failure of the fill material creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Sidecast fill material created from skid trail construction placed on steep slopes can initiate debris slides, torrents or flows.
- Concentrated drainage from skid trails onto unstable areas can initiate debris slides, torrents or flows.
- Cut-slope of skid trails can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of skid trails can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Removal of vegetation from these slopes can result in loss of evapotranspiration and thus increase pore water pressures that could initiate slope failure in this unit.
- Post timber harvest root decay of hardwood or non-redwood conifer species can be a contributing factor in the initiation of debris slides, torrents or flows in this unit.

Confidence:

High confidence for susceptibility of unit to landslides and sediment delivery. Moderate confidence in the placement of this unit. This unit is highly localized and exact boundaries are better determined from field observations. Within this unit there are likely areas of low gradient slopes that are less susceptible to mass wasting.

TSU Number:	3
Description:	Dissected and convergent topography
Materials:	Shallow soils formed from weathered marine sedimentary rocks with localized thin to thick colluvial deposits.
Landforms:	These areas have steep slopes (typically greater than 65%) that have been sculpted over geologic time by repeated debris slide events. The area is characterized primarily by 1) steep convergent and dissected topography located within steep gradient colluvial hollows or headwall swales and small high gradient watercourses, and 2) locally steep planar slopes where there is strong evidence of past landsliding. MRC intends this unit to represent areas with a high hazard potential for shallow landsliding, while not constituting a continuous streamside unit (otherwise it would classify as TSU 1 or 2). The mapped unit may represent isolated individual "high hazard" areas or areas where there is a concentration of "high hazard" areas. Boundaries between higher hazard areas and other more stable areas (i.e. divergent and lower gradient slopes) within the unit should be keyed out as necessary based on field observation of landslide features.
Slope:	Typically >70%, (mean slope of observed mass wasting events is 83%, range is 80%-90%)
Total Area:	506 ac., 5% of the total WAU
MW Processes:	<p>28 <i>road associated landslides</i></p> <ul style="list-style-type: none"> • 28 Debris slides • 0 Debris flows • 0 Debris torrents <p>15 <i>non-road associated landslides</i></p> <ul style="list-style-type: none"> • 13 Debris slides • 1 Debris flows • 1 debris torrents
Non Road-related Landslide Density:	0.03 landslides per acre for the past 62 years.
Forest Practices Sensitivity:	Moderate to high sensitivity to road building, moderate to high sensitivity to harvesting and forest management practices due to moderate to steep slopes within this unit. Localized areas of steeper and/or convergent slopes have even higher sensitivity to forest practices.
Mass Wasting Potential:	High
Delivery Potential:	Moderate

Delivery Criteria

Used:

The converging topography directs mass wasting down slopes toward watercourses. Delivery potential may be high based on relatively high number of debris slides. Landslides in headwater swales often torrent or flow down watercourses. Approximately 80% of landslides in this unit delivered sediment.

Hazard-Potential

Rating:

High

Forest Management

Related Trigger

Mechanisms:

- Sidecast fill material placed on steep slopes can initiate debris slides, torrents or flows in this unit.
- Concentrated drainage from roads onto unstable areas can initiate debris slides, torrents or flows in this unit.
- Concentrated drainage from roads can increase groundwater, accelerating movement of rockslides or earth flows in this unit.
- Poorly sized culvert or excessive debris at watercourse crossings can initiate failure of the fill material creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Concentrated drainage from skid trails onto unstable areas can initiate debris slides, torrents or flows.
- Cut-slope of skid trails can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of skid trails can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Sidecast fill material created from skid trail construction placed on steep slopes can initiate debris slides, torrents or flows.
- Removal of vegetation from these slopes can result in loss of evapotranspiration and thus increase pore water pressures that could initiate slope failure in this unit.
- Post timber harvest root decay of hardwood or non-redwood conifer species can be a contributing factor in the initiation of debris slides, torrents or flows in this unit.

Confidence:

Moderate confidence in placement of unit. This unit is locally variable and exact boundaries are best determined from field observations. Some areas within this unit could have higher susceptibility to landslides and higher delivery rates due to localized areas of steep slopes, weak earth materials, and/or adverse ground water conditions.

TSU Number:	4
Description:	Non-dissected topography
Materials:	Shallow to moderately deep soils formed from weathered marine sedimentary rocks.
Landforms:	Moderate to moderately steep hillslopes with planar, divergent, or broadly convergent slope forms with isolated areas of steep topography or strongly convergent slope forms. TSU 4 is generally a midslope region of lesser slope gradient and more variable slope form than TSU 3.
Slope:	Typically 40% - 65%, (mean slope of observed mass wasting events is 82%, range is 70% - 90%)
Total Area:	7317 acres, 72% of the total WAU
MW Processes:	<p>86 <i>road-associated landslides</i></p> <ul style="list-style-type: none"> • 79 Debris slides • 4 Debris flows • 3 Debris torrents <p>5 <i>non-road associated landslides</i></p> <ul style="list-style-type: none"> • 4 Debris slides • 1 Debris flows • 0 Debris torrent
Non Road-related Landslide Density:	0.0007 landslides per acre for the past 62 years.
Forest Practices Sensitivity:	Moderate sensitivity to road building, moderate to low sensitivity to harvesting and forest management practices due to moderate slope gradients and non-converging topography within this unit. Localized areas of steeper slopes have higher sensitivity to forest practices.
Mass Wasting Potential:	Moderate
Delivery Potential:	High
Delivery Criteria Used:	This unit constitutes a majority of the WAU, which accounts for it having the highest number of landslides. This unit has a low non-road related landslide density, and therefore has a moderate mass wasting hazard. Although landslides in this unit are localized, when landslides occur, the landslide has a high potential to deliver. Approximately 90% of the landslides in this unit delivered sediment. This unit has a moderate sensitivity to road building due to low road landslide density.

Hazard-Potential
Rating:

Moderate

Forest Management
Related Trigger
Mechanisms:

- Sidecast fill material placed on steep slopes can initiate debris slides, torrents or flows in this unit.
- Concentrated drainage from roads onto unstable areas can initiate debris slides, torrents or flows in this unit.
- Concentrated drainage from roads can increase groundwater, accelerating movement of rockslides or earth flows in this unit.
- Poorly sized culvert or excessive debris at watercourse crossings can initiate failure of the fill material creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of roads can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Concentrated drainage from skid trails onto unstable areas can initiate debris slides, torrents or flows.
- Cut-slope of skid trails can expose potential failure planes creating debris slides, torrents or flows in this unit.
- Cut-slope of skid trails can remove support of the toe or expose potential failure planes of rockslides or earth flows.
- Sidecast fill material created from skid trail construction placed on steep slopes can initiate debris slides, torrents or flows.
- Removal of vegetation from these slopes can result in loss of evapotranspiration and thus increase pore water pressures that could initiate slope failure in this unit.
- Post timber harvest root decay of hardwood or non-redwood conifer species can be a contributing factor in the initiation of debris slides, torrents or flows in this unit.

Confidence:

High confidence in placement of unit, however, this unit is locally variable and exact boundaries are best determined from field observations. Some areas within this unit could have higher susceptibility to landslides and higher delivery rates due to localized areas of steep slopes, weak earth materials, and/or adverse ground water conditions.

TSU Number:	5
Description:	Low relief topography
Material:	Moderately deep to deep soils, derived from weathered marine sedimentary rocks.
Landforms:	Characterized by low gradient slopes generally less than 40%, although in some places slopes may be steeper. This unit occurs on ridge crests, low gradient side slopes, and well-developed terraces. Shallow-seated landslides seldom occur and usually do not deliver sediment to stream channels.
Slope:	Typically <40% (based on field observations)
Total Area:	248 acres, 2% of WAU area
MW Processes:	0 <i>landslides</i>
Non Road-related Landslide Density:	0 landslides per acre for past 62 years.
Forest Practices Sensitivity:	Low sensitivity to road building and forest management practices due to low gradient slopes
Mass Wasting Potential:	Low
Delivery Potential:	Low
Delivery Criteria Used:	Sediment delivery in this unit is low.
Hazard-Potential Rating:	Low
Forest Management Related Trigger Mechanisms:	<ul style="list-style-type: none"> • Poorly sized culvert or excessive debris at watercourse crossings can initiate failure of the fill material creating debris slides, torrents or flows in this unit. • Concentrated drainage from roads and skid trails can initiate or accelerate gully erosion, which can increase the potential for mass wasting processes.
Confidence:	High confidence in placement of unit in areas of obviously stable topography. High confidence in mass wasting potential and sediment delivery potential ratings.

Sediment Input from Mass Wasting

Sediment delivery was estimated for shallow-seated landslides in the Rockport Coastal Streams WAU. Depth values were estimated to facilitate approximation of mass for the landslides not observed in the field. In order to extrapolate depth to the shallow-seated landslides not visited in the field, an average was taken from the measured depths of landslides visited in the field. The mean depth of all shallow-seated landslides interpreted as being unrelated to road systems was 4 feet. The mean depth of all shallow seated landslides interpreted as being associated with road systems was also 4 feet. Due to the relative lack of debris flows and torrents, no effort was made to differentiate landslide depths among different shallow landslide types. The mean depth of 4 feet was assigned to all landslides not verified in the field.

The mean sediment delivery percentage assigned to shallow landslides determined to deliver sediment, but not field verified, is 50%. Of the 412 shallow-seated landslides mapped by MRC in this watershed analysis, 385 of the landslides delivered some amount of sediment (Table A-4).

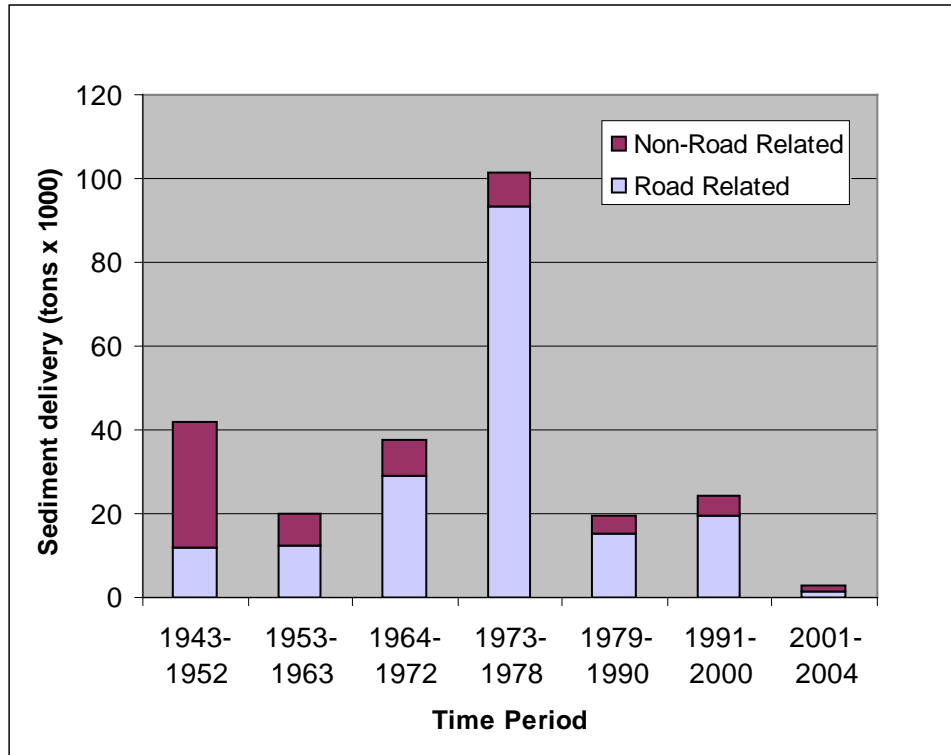
Table A-4. Total Shallow-Seated Landslides Mapped for each PWS in Rockport Coastal Streams WAU.

Planning Watershed	Total Landslides	Landslides with Sediment Delivery	Landslides with No Sediment Delivery
Hardy Creek	99	86	13
Juan Creek	239	228	11
Howard Creek	74	71	3
<i>sum</i>	412	385	27
<i>Percentage</i>	100%	93%	7%

Sediment input to stream channels by mass wasting is quantified for seven time periods (1943-1952, 1953-1963, 1964-1972, 1973-1978, 1979-1990, 1991-2000, 2001-2004). The dates for each of the time periods are based on the date of aerial photographs used to interpret landslides (1952, 1963, 1972, 1978, 1990, 2000, and 2004) and field observations (2006). The available aerial photography did not correspond exactly to ten year time periods for mass wasting assessment, however the time periods and the aerial photographs analyzed approximate decadal intervals and bracket major disturbance events (e.g. intensive tractor logging in the 1960's and 1970's). These time periods allow for a general evaluation of the relative magnitude of sediment delivery rate estimates across the Rockport Coastal Streams WAU.

A total of approximately 246,968 tons of mass wasting sediment delivery was estimated for the time period 1943-2004 in the Rockport Coastal Streams WAU. This equates to approximately 254 tons/sq. mi./yr. Of the total estimated amount, 17% delivered from 1943-1952, 8% delivered from 1953-1963, 15% delivered from 1964-1972, 41% delivered from 1973-1978, 8% delivered from 1979-1990, 10% delivered from 1991-2000, and 1% delivered in the 2001-2004 time period (Chart A-4).

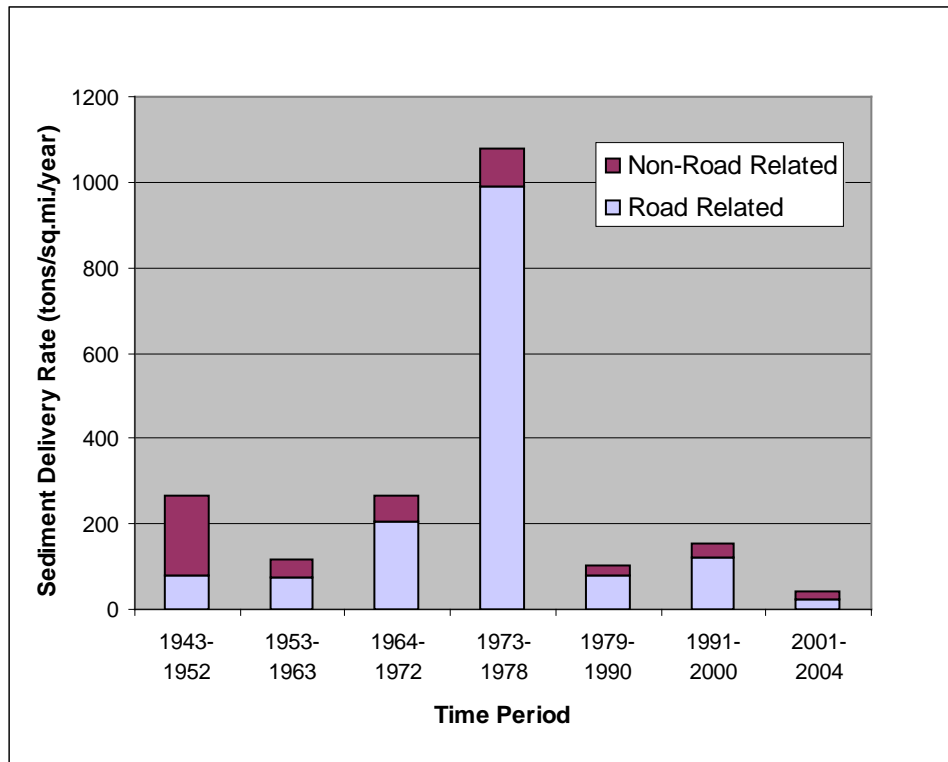
Chart A-4. Sediment Delivery Estimates by Time Period for Shallow-Seated Landslides on MRC Ownership in the Rockport Coastal Streams WAU.



Relatively large amounts of sediment delivered from 1973-1978, particularly in Juan Creek, is the result of intensive road building and ground based yarding on relatively steep slopes. Road and skid trail construction during this era of forest management included the practice of sidecasting excavated fill material, at times onto steep slopes. Additionally, according to local rainfall data, the 1974 storm event produced the wettest days on record at numerous precipitation stations on the northwest coast (Goodridge, 1997). Numerous studies reveal there is a pronounced effect of pore water pressure changes on factor of safety for shallow-seated landslides (Sidle et al., 1985).

The sediment delivery estimates were normalized by time (years) and area (square miles) for the purposes of relative comparison between time intervals and planning watershed. The resulting sediment delivery rates in the Rockport Coastal Streams WAU change dramatically over the time period investigated (Chart A-5).

Chart A-5. Mass Wasting Sediment Delivery Rate (tons/sq.mi./year) from Shallow-Seated Landslides on MRC Ownership in the Rockport Coastal Streams WAU.



Road associated mass wasting (including roads, skid trails, and landings) was found to have contributed 183,044 tons (188 tons/sq. mi./yr) of sediment over the 62 years analyzed in the Rockport Coastal Streams WAU (Table A-5). This represents approximately 74% of the total mass wasting inputs for the Rockport Coastal Streams WAU for 1943-2004. The road related sediment delivery rates for Hardy Creek, Juan Creek, and Howard Creek planning watersheds are quite different. A review of the aerial photo record reveals a majority of Juan Creek had been intensively tractor yarded in the 1970's. Hardy Creek and Howard Creek had not been subjected to the same level of disturbance, this is revealed in the difference in road related sediment rates for the 1973-1978 time period between Hardy Creek (193 tons/sq. mi./yr), Juan Creek (1,806 tons/sq. mi./yr), and Howard Creek (376 tons/sq. mi./yr).

Table A-5. Road Associated Sediment Delivery (tons) for Shallow-Seated Landslides in the Rockport Coastal Streams WAU by Planning Watershed.

Planning Watershed	Road Associated Mass Wasting Sediment Delivery (tons)	Percent of Total Sediment Delivery From Planning Watershed
Hardy Creek	18,612	36%
Juan Creek	134,122	89%
Howard Creek	30,310	67%
<i>Rockport Coastal Streams WAU</i>	<i>183,044</i>	<i>74%</i>

Juan Creek has a higher overall sediment delivery rate from mass wasting than Hardy Creek or Howard Creek over the entire 62 year period (331 tons/sq.mi./yr. in Juan Creek versus 178 tons/sq.mi./yr. in Hardy Creek, and 196 tons/sq.mi./yr. in Howard Creek). The larger sediment delivery rate may be partly due to generally steeper terrain, but is largely attributed to a larger amount of land area disturbed by road construction and tractor logging.

A categorical description of the land area interpreted to concentrate surface and/or subsurface flow to the point of failure for non-road related shallow-seated failures was conducted. Road related failures were excluded because of the many other variables that influence road failures (e.g. thickness of fill, construction techniques, concentrated road run-off, etc.). In this analysis, categories of contributing area included small areas (<0.5 acres), medium sized areas (0.5-3.0 acres) and large areas (>3.0 acres). Areas were determined by a combination of air photo and GIS analysis and indicate a majority of the sediment delivery is occurring from slides where the contributing area is between 0.5 and 3.0 acres in size (Table A-6).

Table A-6. Sediment Delivery from Landslides for MRC Ownership in the Rockport Coastal Streams WAU by Contributing Area.

Planning Watershed	Small Area <0.5 acres	Medium Area 0.5-3.0 acres	Large Area >3.0 acres
Hardy Creek	2,453	19,835	10,525
Juan Creek	750	7,816	7,687
Howard Creek	0	6,480	8,378
<i>Rockport Coastal Streams WAU</i>	<i>3,203</i>	<i>34,131</i>	<i>26,590</i>

Intuitively, a majority of the sediment delivery is occurring from medium and large contributing areas where pore pressure increases in response to precipitation events would be most significant.

A categorical description of the slope aspect for all shallow-seated failures was conducted. Despite the other variables that influence road related failures, as mentioned above, road related failures were included in this analysis. In this analysis slope aspect is determined as an absolute azimuth in the GIS and then categorically described as NE (0°-89°), SE (90°-179°), SW (180°-269°), or NW (270°-359°). Results are presented below (Table A-7).

Table A-7. Sediment Delivery from Landslides for MRC Ownership in the Rockport Coastal Streams WAU by Slope Aspect.

Planning Watershed	NE	NW	SE	SW
Hardy Creek	3,807	11,764	24,464	11,390
Juan Creek	18,415	35,982	45,010	50,968
Howard Creek	1,250	6,188	16,871	20,859
<i>Rockport Coastal Streams WAU</i>	<i>23,472</i>	<i>53,934</i>	<i>86,345</i>	<i>83,217</i>

A majority (69%) of the sediment delivery is occurring on slopes with a predominately south facing aspect. This may be attributed to the south to north direction that rain falls when storm events occur over in the area, resulting in increased pore water pressure increases on south facing slopes.

The distribution of shallow-seated landslides by soil type was analyzed to investigate the relationship between sediment delivery and soil type. The Mendocino County Soil Survey (Rittiman and Thorson, 2001) data includes a classification (USCS, Unified Soil Classification System) that describes the general properties of the soil and allows for a categorical description (Coarse, Fine, or Mixed) based on the distribution of grain size. The GIS was queried for the mapped soil type at the crown of the failure and the USCS soil type was categorically described as either coarse (predominately gravel and sand), fine (predominately silt or clay), or mixed (containing both coarse and fine grain sizes). Criteria for mapping soil types and classifying them based on the USCS are presented elsewhere. A portion of Lower Elk Creek was not made available by previous landowners when soils mapping was conducted, therefore the column “NA” is provided to summarize the amount of sediment that was not classified during this analysis. Results are presented below (Table A-8).

Table A-8. Sediment Delivery from Landslides for MRC Ownership in the Rockport Coastal Streams WAU by Soil Type.

Planning Watershed	Coarse	Fine	Mixed
Hardy Creek	48,337	3,088	0
Juan Creek	139,661	10,714	0
Howard Creek	0	45,168	0
<i>Rockport Coastal Streams WAU</i>	<i>187,998</i>	<i>58,970</i>	<i>0</i>

Results of this analysis reveal a majority of the sediment delivery is occurring from coarse grained soils, however, coarse grained soils also make up a majority of the soils mapped in the Rockport Coastal Streams WAU.

Historically, research on the influence of timber harvesting on slope stability has focused on clear-cutting, or even-aged management, where hydrologic changes are most pronounced. The effect of partial harvest, or uneven-aged management, on slope stability is less well known. This data should not be misinterpreted as present forest conditions on MRC lands have resulted in

a majority of the ownership being in a state of partial harvest. The purpose of this analysis is to begin to generate a long term dataset on the relationship between forest conditions and landslide occurrence. Updates to this watershed analysis over time will build upon this dataset with the intention of identifying any emerging trends in the relationship between forest conditions and sediment delivery from partial harvesting.

The effect that forest stand conditions can have on sediment delivery from shallow-seated landsliding is investigated by attributing recent (2001-2004) non-road related failures with a forest inventory variable titled “structure class.” Stands with similar forest attributes (dominant diameter, dominant vegetation, and canopy density) are described by their structure class as a tool for MRC to assess habitat conditions property wide. Generally, in this process vegetation strata are delineated based on an air photo interpretation of individual similar stands, subsequent field sampling generates empirical information on tree species, diameter, and canopy, and similar strata are grouped together to generate structure classes for habitat description purposes. The findings are summarized below (Table A-9).

Table A-9. Forest Stand Attributes for Recent Non-Road Related Landslides on MRC Ownership in the Rockport Coastal Streams WAU.

Slide ID	Structure Class	Dominant Veg.	Dominant Diameter	Canopy Closure
700	13	Conifer	<8”	<40%
701	10	Conifer/Hardwoods	>8”	>40%
706	22	Conifer	16-24”	>60%
707	22	Conifer	16-24”	>60%
708	10	Conifer/Hardwoods	>8”	>40%

Sediment Input by Terrain Stability Unit

Total mass wasting sediment delivery for the Rockport Coastal Streams WAU was separated into respective Terrain Stability Units. Sediment delivery statistics for each TSU are summarized in Table A-10.

Table A-10. Total Sediment Delivery (in tons) by TSU in the Rockport Coastal Streams WAU (tons)

Sediment Delivery (tons)	TSU						
	1	2	3	4	5	6	8
Road Related	48,371	56,919	22,952	54,802	0	0	0
% of road related	26%	31%	13%	30%	0%	0%	0%
Non-Road Related	41,222	11,702	8,875	2,125	0	0	0
% of non-road related	64%	18%	14%	3%	0%	0%	0%
Total	89,593	68,621	31,827	56,927	0	0	0
% of total delivery	36%	28%	13%	23%	0%	0%	0%
Acres	1,037	1,059	506	7,317	248	0	0
% of WAU area	10%	10%	5%	72%	2%	0%	0%
Ratio- delivery %/area %	3.6	2.7	2.6	0.3	0.0	0.0	0.0

The TSU's with the largest estimated sediment delivery are TSU 1 and 2, which cumulatively are estimated to deliver 64% of the total sediment input for the Rockport Coastal Streams WAU. Combining all high hazard units (TSU 1, 2, and 3) would yield 96% of the estimated non-road related sediment input on approximately 25% of the MRC owned acreage. Combining the moderate and low hazard units (TSU 4, and 5) would yield 4% of the estimated non-road related sediment input off the remaining 75% of the property. One measure of the intensity of mass wasting processes in a given TSU is the amount of sediment produced divided by the area in the TSU. The last row in Table A-10 expresses landslide intensity as the ratio of the percentage of total sediment delivered by the percentage of watershed area in the TSU. A ratio of 1.0 would indicate that the map unit is producing a proportion of the sediment delivery equal to the proportion of the map unit area within the WAU. Values of this ratio greater than 1.0 indicate high landslide rates in a relatively concentrated area. The TSUs with the largest ratios were units 1, 2, and 3, with ratios of 3.6, 2.7, and 2.6, respectively. The smallest ratios are found in units 4 and 5; 0.3 and 0.0, respectively. The ratios suggest that the delineation of the high hazard Terrain Stability Units has captured the majority of the estimated sediment delivery from mass wasting over the past 62 years in the Rockport Coastal Streams WAU.

CONCLUSIONS

In forest environments of the California Coast Range, mass wasting is a common, natural occurrence. In the Rockport Coastal Streams WAU this is due to steep slopes, the condition of weathered and intensely sheared and fractured marine sedimentary rocks, seismic activity, locally thick colluvial soils, a history of timber harvest practices, and the occurrence of high intensity rainfall events. Mass wasting events are episodic and many landslides may happen in a short time frame. Mass wasting features of variable age and stability are observed throughout the Rockport Coastal Streams WAU. A majority of the landslides visited in the field during this assessment occurred on slopes greater than 70%. Seeps and springs were evident in the evacuated cavity at many sites. Particular caution should be exercised when conducting any type of forest management activity in areas with convergent or locally steep topography.

Mass wasting sediment input is estimated to be at least 254 tons/sq.mi./yr. over the 1943-2004 time period for the entire Rockport Coastal Streams WAU. However, approximately 73% of the shallow-seated landslides inventoried in the Rockport Coastal Streams WAU are road associated (includes roads, skid trails, and landings). Road associated mass wasting represented 74% of the estimated sediment delivery, or at least 188 tons/sq. mi./yr of sediment over the 62 years analyzed. Road construction is thus a significant factor in the cause of shallow-seated mass wasting events. Improved road construction practices combined with design upgrades of old roads can reduce anthropogenic sediment input rates and mass wasting hazards.

The steep streamside areas of TSU 1, 2, and 3 contribute the highest amount of the sediment per unit area in the watershed. In the moderate and low hazard units of TSU 4 and 5, a large amount of road associated landslides are occurring, suggesting the need to make improvements on roads within the Rockport Coastal Streams WAU.

REFERENCES

California Department of Conservation, Division of Mines and Geology, 1997. Fault Rupture Hazard Zones in California, Special Publication 42, 47 p.

California Geological Survey, 1999. Factors Affecting Landslides in Forested Terrain. State of California Department of Conservation, Division of Mines and Geology Note 50, 5 p.

Cruden, D.M. and Varnes, D.J. 1996. Landslide types and processes. In: Landslides Investigation and Mitigation, Transportation Research Board, Washington DC, Special Report 247: 36-75.

Dietrich, W.E. and Montgomery, D.R., 1998. SHALSTAB; a digital terrain model for mapping shallow-landslide potential, NCASI Technical Report, February 1998, 29 pp.

Dietrich, W.E., Real de Asua, R., Coyle, J., Orr, B., and Trso, M., 1998. A validation study of the shallow slope stability model, SHALSTAB, in forested lands of Northern California. Stillwater Sciences Internal Report, Berkeley, CA.

Goodridge, J., 1997. California's rainfall records: 1862-1997. CD ROM. CD produced by USDA Forest Service, Pacific Southwest Research Station, Arcata, California.

Keaton, J.R. and DeGraff, J.V., 1996. Surface Observation and Geologic Mapping. In: Landslides Investigation and Mitigation, Transportation Research Board, Washington DC, Special Report 247: 178-230.

Kelley, F.R., 1983. Geology and Geomorphic Features Related to Landsliding, Westport 7.5-Minute Quad, DMG OFR 83-32, Scale 1:24,000.

Kelley, F.R., 1984. Geology and Geomorphic Features Related to Landsliding, Lincoln Ridge 7.5-Minute Quad, DMG OFR 84-14, Scale 1:24,000.

Reeves, G.H., Benda, L.E., Burnett, K.M., Bisson, P.A., Sedell, J.R., 1995. A Disturbance-Based Ecosystem Approach to Maintaining and Restoring Freshwater Habitats of Evolutionary Significant Units of Anadromous Salmonids in the Pacific Northwest. American Fisheries Society Symposium, 17:334-349.

Rittiman, C.A., and Thorson, T., 2001. Soil Survey of Mendocino County, California, Western Part. U.S. Department of Agriculture, Natural Resources Conservation Service, Posted web-site at http://www.ca.nrcs.usda.gov/mlra02/wmendo/eureka_qd.html

Selby, M.J., 1993. Hillslope materials and processes. Second Edition. Oxford University Press. Oxford.

Sidle, R.C., Pearce, A.J., and O'Loughlin, C.L., 1985. Hillslope Stability and Land Use. Water Resources Monograph 11, American Geophysical Union, 140 p.

Swanston, D.N., Lienkaemper, G.W., Mersereau, R.C., and Levno, A.B. 1988. Timber harvest and progressive deformation of slopes in southwestern Oregon. AEG Bulletin, 25(3):371-381.

Washington Forest Practice Board, 1995. Standard methodology for conducting watershed analysis. Version 4.0. WA-DNR Seattle, WA.

Wieczorek, G.F., 1996. Landslide Triggering mechanisms. In: Landslides Investigation and Mitigation, Transportation Research Board, Washington DC, Special Report 247: 76-90.

**Coastal Rockport Mass Wasting Inventory
Appendix A**

Watershed: Rockport Coastal Streams		Shallow-seated landslides																									Deep-seated landslides										Mass Wasting Inventory Sheet							
																																					Mendocino Redwood Company, LLC							
Unique ID#	PWS	T & R	Air Photo year	Air Photo frame	Landslide Type	TSU	Certainty	Length	Width	Depth	Slide Vol.	Sed. Routing	Dom. Aspect	Sed. Del.	Sed. Delivery	Slope	Age	Slope	Slide	Road	Contrib.	Soil	Struc.	Toe	Body	Lat.	Main	DS	Complex	Field	Comments													
	Sec. #				DS DF DT	1 2 3	D P Q	feet	feet	feet	yd ³	P I N	NE SE	25 50 75	yd ³	(%)	N R O	C D P	H S I N	R S L	S M L	USCS	1 to	1 2 3	1 2 3	1 2 3	1 2 3	1 2 3 4	Y N	Y N														
					EF RS	4 5 6						NW SW		100 (%)									24	4 5	4 5	4 5	4 5																	
1	RW	17	1952	4K-163	DT	2	P	100	75	4	1111	P	SW	50	556	750		N	C	S	N	M	F												1000 foot torrent									
2	RW	17	1952	4K-163	DS	1	Q	100	75	4	1111	I	NW	50	556	750		N	P	S	N	L	F																					
3	RW	17	1952	4K-163	DS	3	D	100	50	4	741	I	NE	50	370	500		N	C	H	R	L	F																					
4	RW	9	1952	4K-164	DS	1	P	75	50	4	556	P	SW	50	278	375		N	P	S	N	L	F																					
5	RJ	6	1952	4K-164	DS	1	D	100	75	4	1111	P	SE	50	556	750		R	C	S	R	L	C																					
6	RJ	6	1952	4K-164	DS	1	P	150	75	4	1667	P	SE	50	833	1125		R	C	S	R	L	C																					
7	RJ	5	1952	4K-164	DS	1	P	100	100	4	1481	P	SE	50	741	1000		R	P	S	N	L	C																					
8	RJ	5	1952	4K-164	DF	4	D	75	50	4	556	I	SW	50	278	375		N	C	H	R	L	C																	cutslope failure, 400 foot runout				
9	RJ	4	1952	4K-164	DS	2	D	50	50	4	370	I	SW	50	185	250		N	P	S	N	L	C																					
10	RJ	4	1952	4K-164	DS	2	D	150	125	4	2778	I	SW	50	1389	1875		N	C	S	S	L	C																					
11	RJ	5	1952	4K-165	DS	4	D	100	50	4	741	N	NW	0	0	0		N	C	H	L	L	F																		sidecast fill failure			
12	RJ	6	1952	4K-165	DS	1	D	75	50	4	556	I	SE	50	278	375		N	C	H	N	M	C																					
13	RJ	6	1952	4K-165	DS	1	P	50	50	4	370	I	SE	50	185	250		N	C	S	N	M	C																					
14	RJ	32	1952	4K-165	DS	2	D	100	100	4	1481	I	SW	50	741	1000		N	C	S	N	M	C																					
15	RH	29	1952	4K-166	DS	2	D	150	75	4	1667	P	SE	50	833	1125		N	P	S	N	M	C																					
16	RH	29	1952	4K-166	DS	1	D	200	150	4	4444	P	NW	50	2222	3000		N	C	S	N	L	C																					
17	RH	31	1952	4K-166	DS	2	D	100	50	4	741	P	SW	50	370	500		N	P	S	N	L	C																					
18	RH	30	1952	4K-166	DS	3	D	75	50	4	556	I	SE	50	278	375		N	C	H	N	M	F																					
19	RH	30	1952	4K-166	DS	3	D	100	50	4	741	N	SE	0	0	0		N	C	H	N	M	F																					
20	RH	30	1952	4K-166	DS	3	D	75	50	4	556	I	SW	50	278	375		N	C	H	N	L	F																					
22	RH	29	1952	4K-166	DS	1	D	125	75	3	1042	P	SE	50	521	703		80	N	P	S	N	S	C																			Y	
23	RH	29	1952	4K-166	DS	1	D	100	50	4	741	I	SE	50	370	500			N	C	S	N	M	C																				
24	RH	28	1952	4K-166	DS	3	D	100	50	4	741	I	SW	50	370	500			R	C	H	N	S	C																				
25	RH	32	1952	4K-166	DS	1	D	50	50	4	370	P	SW	25	93	125		70	N	C	S	R	L	C																			possible outslope failure	
26	RH	32	1952	4K-166	DS	1	D	50	50	4	370	P	SW	25	93	125		70	N	C	S	R	L	C																			possible outslope failure	
27	RH	32	1952	4K-166	DS	1	D	75	75	4	833	P	SW	50	417	563			N	C	S	R	L	C																			possible outslope failure	
28	RH	29	1952	4K-166	DS	2	P	100	150	4	2222	N	SW	0	0	0			N	C	N	N	L	C																			possible DSL crown scarp	
29	RH	29	1952	4K-166	DS	1	D	75	50	4	556	P	NE	50	278	375			N	P	S	N	L	C																			possible DSL toe	
30	RH	29	1952	4K-167	DS	3	D	100	100	4	1481	I	NE	50	741	1000			N	C	S	N	M	C																				
31	RH	29	1952	4K-167	DS	3	D	100	100	4	1481	I	NW	50	741	1000			N	C	S	N	M	C																				
32	RH	29	1952	4K-167	DS	3	D	150	50	4	1111	P	SW	50	556	750			O	C	S	N	M	C																				
33	RH	29	1952	4K-167	DF	1	D	100	75	4	1111	I	NW	50	556	750			N	C	H	N	S	F																				250 foot runout
34	RH	20	1952	4K-167	DS	1	D	150	175	6	5833	P	SE	100	5833	7875		100	R	C	S	N	M	C																			Y	
35	RH	21	1952	4K-167	DS	2	D	100	75	4	1111	I	NW	50	556	750			N	C	H	R	L	C																			possible DSL crown scarp	
36	RH	29	1952	4K-167	DS	1	P	50	50	3	278	P	NW	75	208	281		90	N	P	S	N	M	C																			possible DSL crown scarp	
37	RH	28	1952	5K-15	DF	4	D	150	50	4	1111	I	SE	50	556	750			N	C	N	N	L	C																				
38	RH	28	1952	5K-15	DS	1	D	50	50	4	370	P	SE	50	185	250			R	P	S	N	M	C																				
39	RH	28	1952	5K-15	DS	1	D	100	50	4	741	P	SW	50	370	500			O	C	S	N	M	C																				
40	RH	28	1952	5K-15	DS	1	P	75	50	4	556	P	SW	50	278	375			O	C	S	N	M	C																				
41	RH	28	1952	5K-15	DS	2	D	50	50	4	370	P	SE	50	185	250			R	P	S	N	L	C																				
42	RH	28	1952	5K-15	DS	4	D	50	50	4	370	I	NW	50	185	250			N	C	H	N	M	C																				
43	RH	28	1952	5K-15	DS	4	D	50	50	4	370	I	NW	50	185	250			N	C	H	N	M	C																				
44	RJ	34	1952	5K-16	DS	4	D	100	75	4	1111	I	SW	50	556	750			N	C	H	N	S	C																				
45	RJ	4	1952	5K-17	DS	4	D	75	50	3	417	I	SW	25	104	141		90	N	C	H	R	L	C																			Y	
46	RJ	4	1952	5K-17	DS	2	D	100	100	4	1481	I	SW	75	1111	1500			90	N	C	H	R	L	C																		Y	
47	RJ	4	1952	5K-17	DS	4	D	100	75	4	1111	I	NW	75	833	1125			90	N	C	H	L	L	C																		Y	
48	RJ	4	1952	5K-17	DS	2	D	75	40	4	444	I	SE	50	222	300			90	N	C	N	R	L	C																		Y	
49	RJ	4	1952	5K-17	DS	4	D	50	50	4	370	I	SE	50	185	250				N	C	H	R	L	C																			
50	RJ	4	1952	5K-17	DS	2	D	75	50																																			

Watershed: Rockport Coastal Streams										Shallow-seated landslides														Deep-seated landslides										Mass Wasting Inventory Sheet				
Unique ID#	PWS	T & R	Air Photo year	Air Photo frame	Landslide Type	TSU	Certainty	Size			Slide Vol.	Sed. Routing	Dom. Aspect	Sed. Del. Ratio	Sed. Delivery yd³/3	Sed. Delivery tons	Slope (field) (%)	Age	Slope Form	Slide Loc.	Road Assoc.	Contrib. Area	Soil Type	Struc. Class	Toe Activity	Body Morph.	Lat. Scarps	Main Scarps	DS Veg.	Complex	Field Obs.	Comments						
								Length feet	Width feet	Depth feet																							P	I	N	NE	SE	25
110	RH	28	1963	15-133	DF	2	D	50	50	4	370	I	NW	50	185	250		N	C	S	R																	
111	RH	31	1963	15-118	DS	1	D	60	40	3	267	P	SE	50	133	180	75	N	C	S	N	M													Y			
112	RJ	5	1963	15-119	DS	1	P	100	50	4	741	P	SW	50	370	500		N	C	S	N	L	C													possible DSL toe		
113	RJ	32	1963	15-134	DS	4	D	100	75	4	1111	I	SW	50	556	750		N	C	S	R																	
114	RJ	4	1963	15-134	DS	1	D	50	50	4	370	P	SW	50	185	250		N	P	S	R																	
115	RJ	34	1963	15-134	DS	4	D	75	100	4	1111	I	SW	50	556	750		N	C	H	R																	
116	RJ	5	1963	15-134	DS	2	D	75	50	4	556	P	NE	50	278	375		N	P	S	S																	
117	RJ	2	1963	9-60	DS	2	D	50	50	4	370	I	SW	50	185	250		N	P	S	S																	
118	RJ	34	1963	15-134	DS	2	P	50	50	4	370	I	SW	50	185	250		N	C	S	S																	
119	RW	8	1963	15-136	DS	2	D	70	50	4	519	I	SE	25	130	175	70	N	C	H	R																Y	
120	RW	9	1963	15-136	DS	4	D	75	50	4	556	I	SE	50	278	375		O	C	H	R																	
121	RW	9	1963	15-136	DS	1	D	75	50	4	556	P	NW	50	278	375		O	P	S	S																	
122	RW	9	1963	15-136	DS	1	Q	50	50	4	370	P	NW	50	185	250		O	C	S	S																	
123	RW	9	1963	15-136	DS	3	P	200	200	4	5926	P	SW	50	2963	4000		R	C	S	S																	
124	RW	9	1963	15-136	DT	2	D	100	50	4	741	P	SW	50	370	500		R	C	H	R																500 foot torrent track	
125	RW	9	1963	15-136	DS	3	P	150	100	4	2222	P	SE	50	1111	1500		S	P	S	S																	
126	RW	10	1963	15-136	DS	1	D	50	50	4	370	P	SE	50	185	250		N	C	S	S																	
127	RW	9	1963	15-136	DS	2	D	150	150	4	3333	I	SE	50	1667	2250		N	C	H	N	M																possible DSL toe
200	RW	8	1972	5-106	DS	4	D	50	50	4	370	P	NW	50	185	250		R	C	S	N																	
201	RW	8	1972	5-106	DS	1	D	50	75	4	556	P	NW	50	278	375		N	C	S	S																	
202	RW	8	1972	5-106	DS	2	D	100	150	4	2222	I	SW	50	1111	1500		N	C	S	S																	
203	RW	8	1972	5-106	DS	1	D	75	75	4	833	P	NW	50	417	563		N	P	S	R																	
204	RW	8	1972	5-106	DS	1	D	50	50	4	370	P	SW	50	185	250		N	P	S	N	L																
205	RW	16	1972	5-155	DS	2	P	50	50	4	370	P	NW	50	185	250		R	C	S	N	L																
206	RW	16	1972	5-155	DS	3	Q	50	50	4	370	P	SW	50	185	250		N	P	N	S																	
207	RW	10	1972	5-155	DS	1	D	100	75	4	1111	P	SW	50	556	750		N	C	S	N	M																
208	RW	10	1972	5-155	DS	4	D	100	100	4	1481	P	NW	50	741	1000		N	P	S	R																	cutslope failure
209	RJ	6	1972	5-108	DS	1	D	50	50	3	278	P	NW	50	139	188	80	N	C	S	R																Y	
210	RJ	6	1972	5-108	DS	1	D	75	100	4	1111	P	SE	50	556	750		N	C	S	S																	
211	RJ	6	1972	5-108	DS	1	D	75	100	4	1111	P	SE	25	278	375	75	N	P	N	R																Y	
212	RJ	6	1972	5-108	DT	1	D	50	60	4	444	I	SE	100	444	600	80	N	C	S	R																Y	
213	RJ	6	1972	5-108	DS	2	D	75	75	5	1042	I	SE	50	521	703	80	N	C	S	R																Y	
214	RJ	6	1972	5-108	DS	1	D	100	100	4	1481	P	SE	50	741	1000		R	C	S	R																	
215	RJ	6	1972	5-108	DF	4	D	75	50	4	556	P	NW	50	278	375		N	P	S	R																	300 foot run-out outside meander bend
216	RJ	6	1972	5-108	DS	1	D	100	50	4	741	P	SE	50	370	500		N	P	I	N	L																
217	RJ	5	1972	5-108	DS	2	D	50	50	4	370	I	SW	50	185	250		N	C	S	R																	
218	RJ	6	1972	5-108	DS	3	P	200	200	4	5926	P	SW	50	2963	4000		R	C	S	R																	possible DSL
219	RJ	5	1972	5-108	DS	1	D	50	50	4	370	P	SW	50	185	250		N	P	S	R																	
220	RJ	5	1972	5-108	DS	1	P	50	50	4	370	P	SW	50	185	250		N	P	S	R																	
221	RJ	5	1972	5-108	DS	1	D	50	70	5	648	P	SE	75	486	656	70	N	P	S	R																Y	
222	RJ	5	1972	5-108	DS	1	D	40	40	4	237	P	SE	100	237	320	70	N	P	S	R																Y	
223	RJ	6	1972	5-108	DS	1	D	100	100	4	1481	P	SW	50	741	1000		N	C	S	R																	cutslope failure
224	RJ	32	1972	5-108	DS	2	D	100	50	4	741	P	SW	50	370	500		N	P	S	L																	
225	RJ	4	1972	5-151	DS	1	D	150	50	4	1111	P	NW	50	556	750		N	P	S	N	M																
226	RJ	4	1972	5-151	DT	2	D	150	150	4	3333	P	NE	50	1667	2250		N	C	S	R																	500 foot torrent track
227	RJ	4	1972	5-																																		

Watershed: Rockport Coastal Streams				Shallow-seated landslides																	Deep-seated landslides									Mass Wasting Inventory Sheet						
																														Mendocino Redwood Company, LLC						
Unique PWS	T & R	Air Photo	Air Photo	Landslide	TSU	Certainty	Size		Slide	Sed.	Dom.	Sed. Del.	Sed.	Sed.	Slope	Age	Slope	Slide	Road	Contrib.	Soil	Struc.	Toe	Body	Lat.	Main	DS	Complex	Field	Comments						
#	Sec. #	year	frame	Type			Length	Width	Depth	Vol.	Routing	Aspect	Ratio	Delivery	Delivery	(field)	N R O	C D P	H S I N	R S L	S M L	USCS	1 to	1 2 3	1 2 3	1 2 3	1 2 3	1 2 3 4	Y N	Y N						
				DS DF DT	1 2 3	D P Q	feet	feet	feet	yd ³	P I N	NE SE	25 50 75	yd ³	tons	(%)																				
				EF RS	4 5 6						NW SW	100 (%)										24	4 5	4 5	4 5	4 5										
414	RJ	4	1978	3-4	DS	2	D	75	30	3	250	I	SE	50	125	169	90	N	C	N	S										Y					
415	RJ	4	1978	3-4	DS	4	D	75	50	4	556	I	NE	50	278	375		N	P	H	S															
416	RJ	4	1978	3-4	DS	2	D	75	50	4	556	I	NE	50	278	375		N	P	S	R															
417	RJ	4	1978	3-4	DS	2	D	75	50	4	556	I	NW	50	278	375		N	P	N	R															
418	RJ	4	1978	3-4	DS	2	D	50	50	4	370	I	NE	50	185	250		N	P	N	R															
419	RH	31	1978	2-4	DS	4	D	100	50	4	741	P	SW	50	370	500		N	P	S	R															
420	RH	31	1978	2-4	DS	1	D	75	45	4	500	I	SE	50	250	338	70	N	C	S	R												Y			
421	RH	31	1978	2-4	DS	1	D	75	75	4	833	P	NW	50	417	563		N	P	S	N	M														
422	RH	32	1978	3-4	DS	1	D	100	80	4	1185	P	SE	50	593	800	80	N	C	S	R												Y	cutslope failure		
423	RH	32	1978	3-4	DS	1	D	45	40	4	267	P	NW	50	133	180	75	N	P	S	N	M											Y			
424	RH	32	1978	3-4	DS	1	D	100	75	4	1111	P	NE	50	556	750		N	P	S	N	L														
425	RH	32	1978	3-4	DS	1	D	100	75	4	1111	P	SW	50	556	750		N	C	S	R															
426	RH	33	1978	3-4	DS	4	Q	75	50	4	556	I	NW	50	278	375		N	C	S	S															
427	RH	29	1978	3-6	DS	1	P	75	50	4	556	P	NW	50	278	375		N	P	S	N	L														
428	RH	29	1978	3-6	DS	1	P	75	50	4	556	P	NW	50	278	375		N	P	S	N	L														
429	RH	20	1978	3-6	DS	4	D	50	50	4	370	P	SE	50	185	250		N	C	S	S															
430	RH	20	1978	3-6	DS	1	D	75	50	4	556	P	SW	50	278	375	80	N	P	S	R													Y		
431	RH	20	1978	3-6	DS	1	D	75	75	3	625	I	SE	75	469	633	85	N	P	S	R													Y		
432	RH	21	1978	3-6	DS	4	D	100	75	4	1111	I	SW	50	556	750		N	P	S	S															
433	RH	28	1978	3-6	DS	4	P	75	50	4	556	I	SE	50	278	375		N	C	H	S															
434	RH	28	1978	3-6	DS	3	P	50	50	4	370	I	SE	50	185	250		N	C	H	S															
435	RJ	3	1978	4-4	DS	4	D	150	75	4	1667	I	SW	50	833	1125		N	C	S	S															
436	RJ	3	1978	4-4	DS	2	D	150	50	4	1111	I	NW	50	556	750		N	C	S	S															
437	RJ	2	1978	4-4	DS	1	P	50	50	4	370	P	SW	50	185	250		N	P	S	S															
438	RJ	2	1978	4-4	DS	4	P	75	50	4	556	P	SW	50	278	375		N	C	S	S															
439	RJ	34	1978	4-6	DS	2	P	100	100	4	1481	I	SW	50	741	1000		N	P	S	S															
440	RJ	2	1978	5-4	DS	2	D	75	75	4	833	I	SW	50	417	563		N	C	S	S															
441	RJ	2	1978	5-4	DS	4	D	100	50	4	741	I	SW	50	370	500		N	P	S	S															
442	RJ	2	1978	5-4	DS	3	D	75	50	4	556	I	SW	50	278	375		N	C	S	S															
500	RW	8	1990	M4-4	DS	2	D	50	50	4	370	I	SW	50	185	250		N	C	S	R															
501	RW	8	1990	M4-4	DS	2	D	50	50	4	370	I	NW	25	93	125	75	N	C	S	R														Y	
502	RW	8	1990	M4-4	DS	4	D	150	150	4	3333	N	SW	0	0	0	75	N	P	N	R														Y	
503	RW	8	1990	M4-4	DS	1	D	50	50	4	370	I	SE	50	185	250		N	P	S	N	M														
504	RW	8	1990	M4-4	DS	1	D	50	50	4	370	I	SE	50	185	250		N	P	S	N	M														
505	RW	9	1990	M4-4	DS	2	D	70	40	3	311	I	SW	75	233	315	80	N	P	S	N	L													Y	
506	RW	9	1990	M5-5	DS	2	D	75	50	4	556	I	SE	50	278	375		R	C	S	R															
507	RJ	8	1990	M3-2	DS	2	D	60	50	3	333	N	NW	0	0	0	80	N	C	H	R														Y	
508	RW	9	1990	M5-3	DS	1	D	50	25	4	185	P	SE	50	93	125		N	C	S	N	L														
509	RW	9	1990	M5-3	DS	4	D	50	25	4	185	N	SW	0	0	0		N	P	H	R															
510	RW	9	1990	M5-5	DS	4	D	150	50	4	1111	I	SW	50	556	750		N	P	N	R															
511	RW	3	1990	M5-5	DS	4	D	75	50	4	556	N	SW	0	0	0	75	N	P	H	R														Y	
512	RW	10	1990	M5-5	DS	1	D	50	50	4	370	I	NW	50	185	250		N	P	S	S															
513	RW	10	1990	M5-5	DS	2	P	50	50	4	370	I	SE	50	185	250		N	C	S	S															
514	RW	10	1990	M5-5	DS	3	P	50	50	4	370	I	NW	50	185	250		N	C	H	S															
515	RW	10	1990	M6-2	DS	2	P	100	50	4	741	I	SW	50	370	500		N	C	S	S															
516	RW	10	1990	M6-2	DS	2	Q	75	50	4	556	I	SW	50	278	375		N	C	S	S															
517	RJ	7	1990	M3-2	DS	2	D	75	50	4	556	I	NW	50	278	375		N	C	S	R															
518	RJ	6	1990	M3-4	DS	2	D	50	25	4	185	I	SW	50	93	125		N	P	S	N	L														
519	RJ	6	1990	M3-4	DS	4	D	50	50	4	370	N	NW	0	0	0	85	N	C	N	R															Y
520	RJ	6	1990	M3-4	DS	1	D	50	50	4	370	P	SW	50	185	250		N	C	S	N	L														
521	RJ	5	1990	M3-4	DS	4	D	75	45	3	375	I	SE	25	94	127	90	N	P	N	R														Y	
522	RJ	32	1990	M3-4	DS	4	D	75	50	3	417	I	SE	25	104	141	90	N	P	N	R														Y	
523	RJ	5	1990	M4-4	DS	4	D	75	50	4	556	I	NW	50	278	375		N	P	S	R															
524	RJ	5	1990	M4-4	DS	1	D	100	100	4	1481	P	SE	50	741	1000		R	P	N	S															
525	RJ	5	1990	M4-4	DS	4	D	100	50	4	741	P	SW	50	370	500		R	P	N	R															
526	RJ	5	1990	M4-4	DS	3	D																													

Watershed: Rockport Coastal Streams		Shallow-seated landslides													Deep-seated landslides										Mass Wasting Inventory Sheet Mendocino Redwood Company, LLC															
Unique #	PWS Sec. #	T & R year	Air Photo frame	Air Photo frame	Landslide Type	TSU	Certainty	Length feet	Width feet	Depth feet	Slide Vol. yd ³	Sed. Routing	Dom. Aspect	Sed. Del. Ratio	Sed. Delivery yd ³	Sed. Delivery tons	Slope (field)	Age	Slope Form	Slide Loc.	Road Assoc.	Contrib. Area	Soil Type	Struct. Class	Toe Activity	Body Morph.	Lat. Scarps	Main Scarps	DS Veg.	Complex	Field Obs.	Comments								
					DS DF DT EF RS	1 2 3 4 5 6	D P Q					P I N	NE SE NW SW	25 50 75 100 (%)			(%)	N R O	C D P	H S I N	R S L	S M L	USCS	1 to 24	1 2 3 4 5	1 2 3 4 5	1 2 3 4 5	1 2 3 4 5	1 2 3 4 Y N											
536	RJ	33	1990	M4-6	DS	2	P	50	25	4	185	N	SW	0	0	0		N	P	N	N	M	C																	
537	RJ	33	1990	M4-6	DS	2	D	100	50	4	741	I	NW	50	370	500		N	P	N	S		C																	
538	RJ	4	1990	M5-7	DS	2	D	125	50	4	926	N	NW	0	0	0		N	C	N	R		C																	
539	RJ	34	1990	M5-7	DS	1	P	50	75	4	556	P	NW	50	278	375		N	P	S	S		C																	
540	RJ	34	1990	M5-7	DS	1	P	75	75	4	833	P	SE	50	417	563		R	P	S	R		C																	
541	RJ	3	1990	M6-4	DS	4	D	50	50	4	370	N	NW	0	0	0		N	C	H	R		C																	
542	RJ	2	1990	M6-4	DS	1	D	50	50	3	278	P	NW	75	208	281	75	N	P	S	R		C												Y					
543	RJ	2	1990	M6-4	DS	1	D	60	40	4	356	P	SE	25	89	120	80	N	P	S	R		C													Y				
544	RJ	3	1990	M6-4	DS	2	P	50	50	4	370	I	NW	50	185	250		N	P	S	S		C																	
545	RJ	34	1990	M6-6	DS	1	D	75	50	4	556	I	SW	50	278	375		N	P	S	S		C																	
546	RJ	34	1990	M6-6	DS	1	P	75	50	4	556	I	SE	50	278	375		N	C	S	S		C																	
547	RJ	34	1990	M6-6	DS	2	D	100	50	4	741	I	NW	75	556	750	75	N	C	S	S		C													Y	200 foot runout, enlargement of #394			
548	RJ	2	1990	M7-2	DS	1	D	100	150	4	2222	P	SW	50	1111	1500		N	P	S	S		C																	
549	RH	36	1990	M2-3	DS	1	D	125	75	4	1389	P	SW	0	0	0		N	P	N	N	M	F																	
550	RH	31	1990	M3-6	DS	2	D	40	35	4	207	I	SW	25	52	70	75	N	C	S	R		F													Y				
551	RH	31	1990	M3-6	DS	1	D	50	50	3	278	P	NE	75	208	281	85	N	C	S	N	M	C															Y		
552	RH	33	1990	M4-6	DS	1	D	100	50	4	741	P	SE	50	370	500		N	P	S	N	L	C																	
553	RH	33	1990	M4-6	DS	1	D	100	75	4	1111	P	SE	50	556	750		N	P	S	N	L	C																	
554	RH	29	1990	M4-8	DS	4	D	50	25	4	185	I	SE	50	93	125		N	P	S	N	M	C																	
555	RH	29	1990	M4-8	DS	2	P	75	75	4	833	I	SE	25	208	281	120	N	C	S	S		C															Y		
556	RH	29	1990	M4-8	DS	2	Q	50	50	3	278	I	SW	25	69	94	90	N	C	S	S		C															Y		
557	RH	29	1990	M4-8	DS	2	P	50	50	2	185	I	SW	100	185	250	85	N	P	S	S		C																Y	
558	RH	28	1990	M4-8	DS	3	D	75	50	4	556	N	SE	0	0	0		N	C	H	R		C																	
559	RH	29	1990	M4-8	DS	3	D	150	100	4	2222	N	NW	0	0	0		N	C	N	R		C																	
560	RH	28	1990	M5-9	DS	4	D	75	50	4	556	N	SW	0	0	0		N	C	N	R		C																	
561	RH	28	1990	M5-9	DS	4	D	75	50	4	556	N	NW	0	0	0		N	P	N	R		C																	
562	RH	28	1990	M5-9	DS	2	D	50	25	4	185	N	SW	0	0	0		N	P	N	S		C																	
563	RH	28	1990	M5-9	DS	2	D	75	75	4	833	I	SW	50	417	563		N	C	H	R		C																	
564	RH	21	1990	M5-9	DS	2	Q	50	50	4	370	I	SE	50	185	250		N	P	S	S		C																	
565	RH	29	field obs		DS	2	D	75	50	3	417	I	SE	75	313	422	80	O	P	S	N	M	C																Y	
566	RH	29	field obs		DS	1	D	50	40	3	222	I	NW	25	56	75	90	O	P	S	N	M	C																	Y
600	RW	8	2000	3-2	DS	3	Q	150	75	4	1667	I	SW	50	833	1125		N	C	H	N	L	F																	
601	RW	8	2000	3-2	DS	2	D	50	50	3	278	I	SE	25	69	94	80	N	C	S	R		F																	
602	RW	9	2000	4-2	DS	1	D	150	125	4	2778	P	SE	50	1389	1875		R	P	S	R		F																	
603	RW	17	2000	4-2	DS	1	D	75	50	4	556	P	SE	50	278	375		R	C	S	R		F																	
604	RW	8	2000	4-2	DS	3	D	100	75	4	1111	I	NE	50	556	750		N	C	H	N	M	F																	
605	RW	8	2000	4-2	DS	3	P	50	50	4	370	I	SE	50	185	250		N	P	N	N	M	F																	
606	RW	9	2000	5-4	DS	1	D	50	50	4	370	P	SW	50	185	250		N	C	S	N	M	F																	
607	RW	9	2000	5-4	DS	1	D	75	50	4	556	P	SE	50	278	375		N	C	S	S		F																	
608	RW	10	2000	5-4	DS	1	D	150	75	4	1667	P	SE	50	833	1125		N	P	S	N	L	F																	
609	RW	9	2000	5-4	DS	1	P	50	50	4	370	P	SE	50	185	250		N	P	S	N	L	F																	
610	RW	9	2000	5-4	DF	4	D	75	45	5	625	I	SE	75	469	633	75	N	C	H	R		F																	
611	RW	9	2000	5-4	DS	2	D	75	50	4	556	I	SE	25	139	188	80	N	C	H	R		F																	
612	RW	16	2000	6-2	DS	1	P	100	75	4	1111	P	SW	50	556	750		N	C	S	R		F																	
613	RW	10	2000	6-2	DT	2	D	100	75	4	1111	P	SE	50	556	750		N	C	H	R		F																	
614	RW	10	2000	6-2	DS	4	D	5																																

Watershed: Rockport Coastal Streams										Shallow-seated landslides															Deep-seated landslides										Mass Wasting Inventory Sheet	
Unique ID#	PWS	T & R	Air Photo year	Air Photo frame	Landslide Type	TSU	Certainty	Size			Slide Vol. yd³	Sed. Routing	Dom. Aspect	Sed. Del. Ratio	Sed. Del. Delivery yd³	Sed. Del. Delivery tons	Slope (field) (%)	Age	Slope Form	Slide Loc.	Road Assoc.	Contrib. Area	Soil Type	Struc. Class	Toe Activity	Body Morph.	Lat. Scarps	Main Scarps	DS Veg.	Complex	Field Obs.	Comments				
								Length feet	Width feet	Depth feet																							P	I	N	NE
634	RH	36	2000	2-1	DS	1	D	150	50	4	1111	N	SW	0	0	0		N	C	N	N	M	F									coastal bluff retreat				
635	RH	36	2000	2-1	DS	1	D	125	50	4	926	N	SW	0	0	0		N	C	N	N	M	F									coastal bluff retreat				
636	RH	36	2000	2-1	DS	1	D	100	75	4	1111	N	SW	0	0	0		N	C	N	N	M	F									coastal bluff retreat				
637	RH	31	2000	3-4	DT	3	D	100	50	4	741	P	NE	50	370	500		N	C	H	N	S	F									500 foot torrent track				
638	RH	31	2000	3-4	DS	1	D	175	100	7	4537	P	SE	25	1134	1531	80	N	P	S	R	C							Y		cutslope failure overtopped road					
639	RH	29	2000	4-8	DS	1	Q	75	100	4	1111	P	SE	50	556	750		N	C	S	R	C									uncertain about road association					
640	RH	28	2000	5-8	DS	2	D	100	100	4	1481	P	SE	50	741	1000		R	C	S	L	C														
641	RH	28	2000	5-8	DT	4	D	150	75	4	1667	P	SW	50	833	1125		N	C	H	R	C										800 foot torrent track				
642	RH	29	2000	5-8	DS	4	D	75	50	4	556	N	NE	0	0	0		N	C	N	L	F														
643	RW	8	2000	4-3	DS	4	D	150	75	5	2083	I	SE	50	1042	1406	75	N	C	H	R	F										Y				
644	RH	21	2000	5-8	DT	4	D	100	50	4	741	I	SW	50	370	500		N	C	H	R	C										500 foot torrent track				
645	RJ	6	2000	3-2	DS	2	D	50	35	4	259	I	NE	25	65	88	70	R	P	N	R	F										Y				
700	RJ	9	2004	11-68	DS	2	Q	50	50	4	370	I	NE	50	185	250		N	C	S	N	L	C	13												
701	RJ	32	2004	11-70	DS	3	D	100	50	4	741	I	SE	50	370	500		N	C	H	N	M	C	10												
702	RJ	3	2004	13-168	DS	2	P	50	50	4	370	I	NW	50	185	250		N	P	S	S	C														
703	RJ	3	2004	13-170	DS	1	D	75	50	4	556	I	NE	50	278	375		N	P	S	R	C											cutslope failure, delivery uncertain			
704	RJ	34	2004	13-170	DS	4	D	100	50	4	741	I	SE	50	370	500		N	P	N	R	C											cutslope failure, delivery uncertain			
705	RH	28	2004	12-126	DS	4	D	100	50	4	741	N	NW	0	0	0		N	C	H	R	C														
706	RH	32	field obs		DS	1	D	35	40	2	104	P	NW	100	104	140	105	N	P	S	N	M	C	22									Y			
707	RH	20	field obs		DS	1	D	75	40	3	333	P	SW	50	167	225	110	N	P	S	N	L	C	22									Y			
708	RJ	3	field obs		DS	1	D	35	25	3	97	P	NE	100	97	131	80	N	P	I	N	L	C	10									Y			
709	RJ	33	field obs		DS	2	D	40	40	5	296	P	SE	100	296	400	55	N	P	N	R	C											Y			
800			2000	4-2	RS		P	500	300															3	3	3	3	4	N							
801			2000	4-2	RS		Q	750	500															4	4	4	3	4	N							
802			2000	4-2	RS		Q	600	500															3	4	4	3	4	N							
803			2000	4-2	RS		P	1000	750															2	3	3	4	4	N							
804			2000	4-2	RS		P	600	900															3	3	3	2	4	Y							
805			2000	5-3	RS		D	500	500															3	3	2	3	4	N							
806			2000	5-3	RS		P	2000	750															3	3	3	3	4	N							
807			2000	5-3	RS		P	1300	1000															2	3	3	3	4	N							
808			2000	5-3	RS		P	750	500															2	3	3	3	4	N							
809			2000	5-3	RS		Q	2300	1000															3	2	3	3	4	Y							
810			2000	5-3	RS		Q	1200	500															2	3	3	3	4	N							
811			2000	6-2	RS		D	1500	400															3	3	2	2	4	N							
812			2000	6-2	RS		P	1200	400															3	4	3	3	4	N							
813			2000	6-2	RS		P	1500	750															3	2	3	3	4	Y							
814			2000	6-4	RS		P	1500	1200															4	3	3	2	4	Y							
815			2000	6-2	RS		P	700	500															3	3	2	2	3	N							
816			2000	3-4	RS		P	1100	500															3	4	4	4	4	N							
817			2000	3-4	RS		Q	1500	700															3	2	4	4	4	N							
818			2000	3-4	RS		P	1100	500															2	2	3	2	4	N							
819			2000	3-4	RS		Q	1700	1000															2	2	3	2	4	N							
820			2000	4-4	RS		P	2000	700															4	3	4	3	4	N							
821			2000	4-4	RS		P	1100	500															3	3	4	3	4	N							
822			2000	4-6	RS		P	1100	500															4	3	3	3	4	N							
823			2000	5-4	RS		P	700	400															3	4	4	4	4	N							
824			2000	5-6	RS		D	800	500															2	3	2	2	4	N							
825			2000	5-6	RS		D	2000	1400															2	2	2	2	4	N							
826			2000	5-6	RS		P	1000	500															2	2	3	2	4	N							
827			2000	5-6	RS		P	700	400															3	2	2	2	4	N							
828			2000	5-6	RS		D	1800	800															3	2	2	2	4	N							
829			2000	5-6	RS		P	800	800															2	3	3	3	4	N							
830			2000	6-4	RS		P	2000	1100															3	2	3	3	4	N							
831			2000	6-4	RS		P	1000	500															2	3	4	3	4	N							
832																																				

**Rockport Coastal Streams
Watershed Analysis
Unit**

**Map A-1
Mass Wasting Inventory**

This map presents the location of mass wasting features identified on the MRC land in the Rockport Coastal Streams watershed. The mass wasting features were developed from an interpretation of aerial photographs from the 1950s-2004 with field observations taken in 2006. All shallow-seated landslides are identified as a point plotted on the map at the interpreted head scarp of the failure. Deep-seated landslides are represented as a polygon representing the interpreted perimeter of the landslide feature. Physical and geomorphic characteristics of mapped landslides are categorized in a database in the mass wasting report for the Rockport Coastal Streams WAU (Section A).

Shallow-Seated Slide

Volume (cubic yards)

- < 500
- 500 - 5,000
- > 5,000

▣ Deep-Seated Slide

■ MRC Ownership

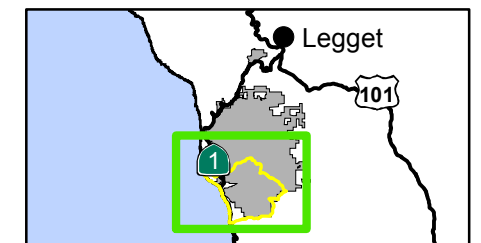
— Planning Watershed Boundary

— Rockport Coastal Streams
Watershed Analysis Unit Boundary

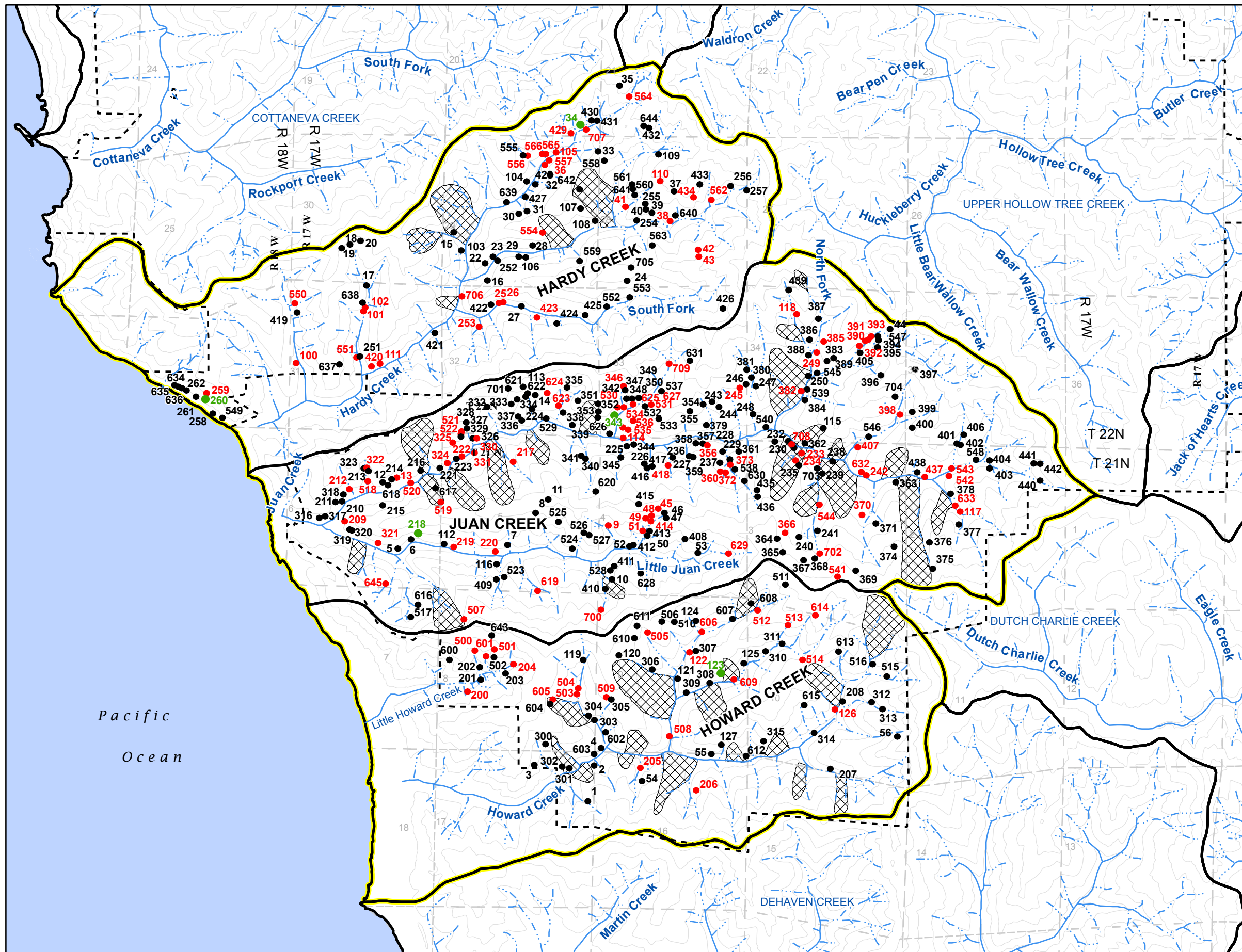
— 200' Contour Interval

Flow Class

- Class I
- Class II
- Class III



0 0.5 1 Miles



**Rockport Coastal Streams
Watershed Analysis
Unit**

**Map A-2
Terrain Stability Units**

This map presents an interpretation of the terrain stability units (TSUs) delineated for the Rockport Coastal Streams WAU. The TSUs characterize the landscape by similar geomorphic attributes, shallow-seated landslide potential, and sediment delivery potential. The TSU designations for the Rockport Coastal Streams WAU are only meant to be general characterizations of similar geomorphic and terrain characteristics related to shallow-seated landslides. Deep-seated landslides are also shown on this map. The deep-seated landslides have been included to provide land managers with supplemental information to guide evaluation of harvest planning and subsequent needs for geologic review. The landscape and geomorphic setting in the Rockport Coastal Streams WAU is certainly more complex than generalized TSUs delineated for this evaluation. The TSUs are only meant to be a starting point for gauging the need for site-specific field assessments. Field observations will override unit boundaries of this map.

Terrain Stability Units

Unit	Description
1	Steep slopes along low gradient watercourses
2	Steep slopes adjacent to select intermittent or ephemeral streams
3	Steep dissected or convergent topography
4	Non-dissected topography
5	Low relief topography
6	Identified earthflow complexes
8	Ohlson Ranch Formation

Deep-Seated Slide

MRC Ownership

Planning Watershed Boundary

Rockport Coastal Streams Watershed Analysis Unit Boundary

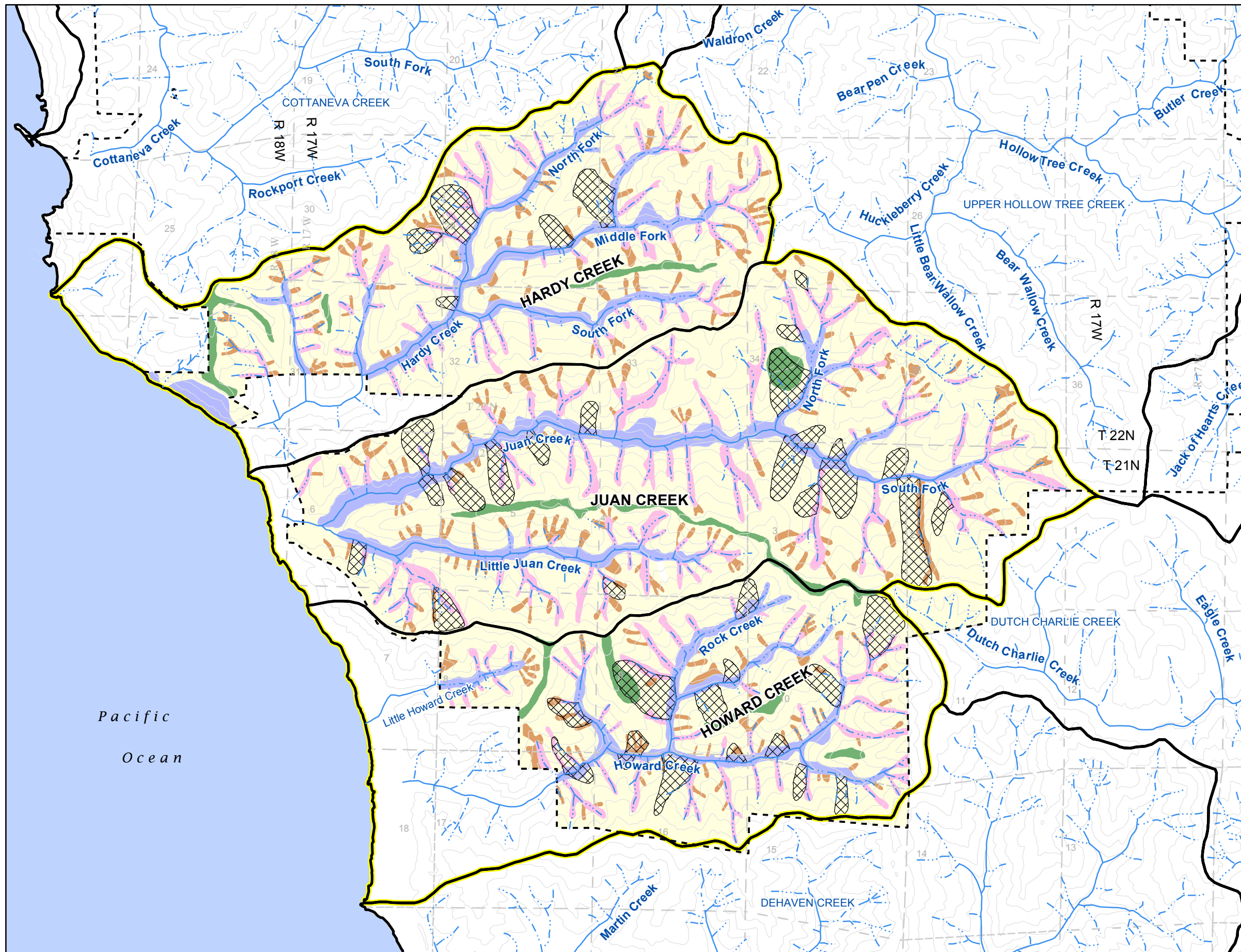
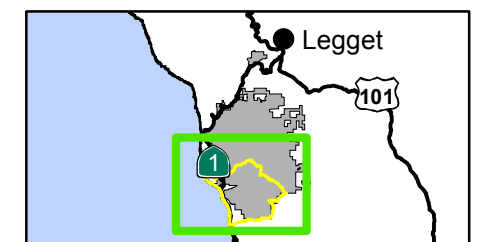
200' Contour Interval

Flow Class

Class I

Class II

Class III



Pacific
Ocean